



## ***Keynote lecture***

**Using connectivity to assess soil erosion and mass movement processes in the landscape:  
applications and discussion of a new paradigm**



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***CONNECTEUR – Scientific Kickoff Meeting (COST Action ES 1306)  
Wageningen, The Netherlands, on August 24-26, 2014***

 National Research Council of Italy





# Why a need for Connectivity paradigm and modelling efforts?

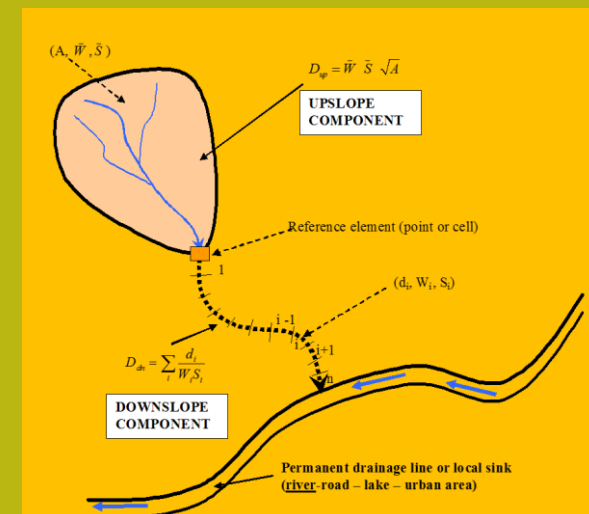


*Northern Apennines (Samoggia valley), ITALY,  
photo Borselli, autumn 2001*

**Flow Connectivity Approach (FCA)**

# Key points of this presentation :

- Connectivity basic concepts (my personal view)
- Development of index of connectivity (IC)
- Theoretical basis of IC index
- Examples of application
- Calibration of IC index and IC interpretation
- Second group of application examples
- Variants of IC index: applications and extension to mass movements
- Discussion on IC index, presents limits and possible future developments
- Highlights and Speculations





*Hydrological connectivity is a term often used to describe the internal linkages between runoff and sediment sources in upper parts of catchments and the corresponding sinks (Croke et al., 2005).*



Tuscany, Italy  
2006

*Connectivity paradigm has prominent importance also in geomorphology and **landscape evolution**: processes and rates...*



Colima, Mexico,  
2011

**SOURCE**



**MASS and Energy  
TRANSFER:  
WATER , SEDIMENTS,....**

**SINK**

( see definitions review in Bracken et al. (2013) )



# Definition of connectivity for sediment flow:

Connectivity may be defined as the chances that a particle has to reach the nearest sink and it depends on: distance to the sink; characteristics of the route; water available to transport from upslope; water that is gained/lost along the downslope route



Erosion and connectivity in an old *biancana badland* levelled field (Tuscany, Italy 2001)

Diffuse connectivity it is also influenced (Cammerrat ,2002) by :

- 1) soil surface irregularity (roughness), which could be very low at the patch scale, but higher at the hillslope and the catchment scales;
- 2) spatial organization of the vegetation at the hillslope scale and the spatial arrangement between land units at the catchment scale;
- 3) rainfall intensity, event duration, and thus the effective rainfall!!!!

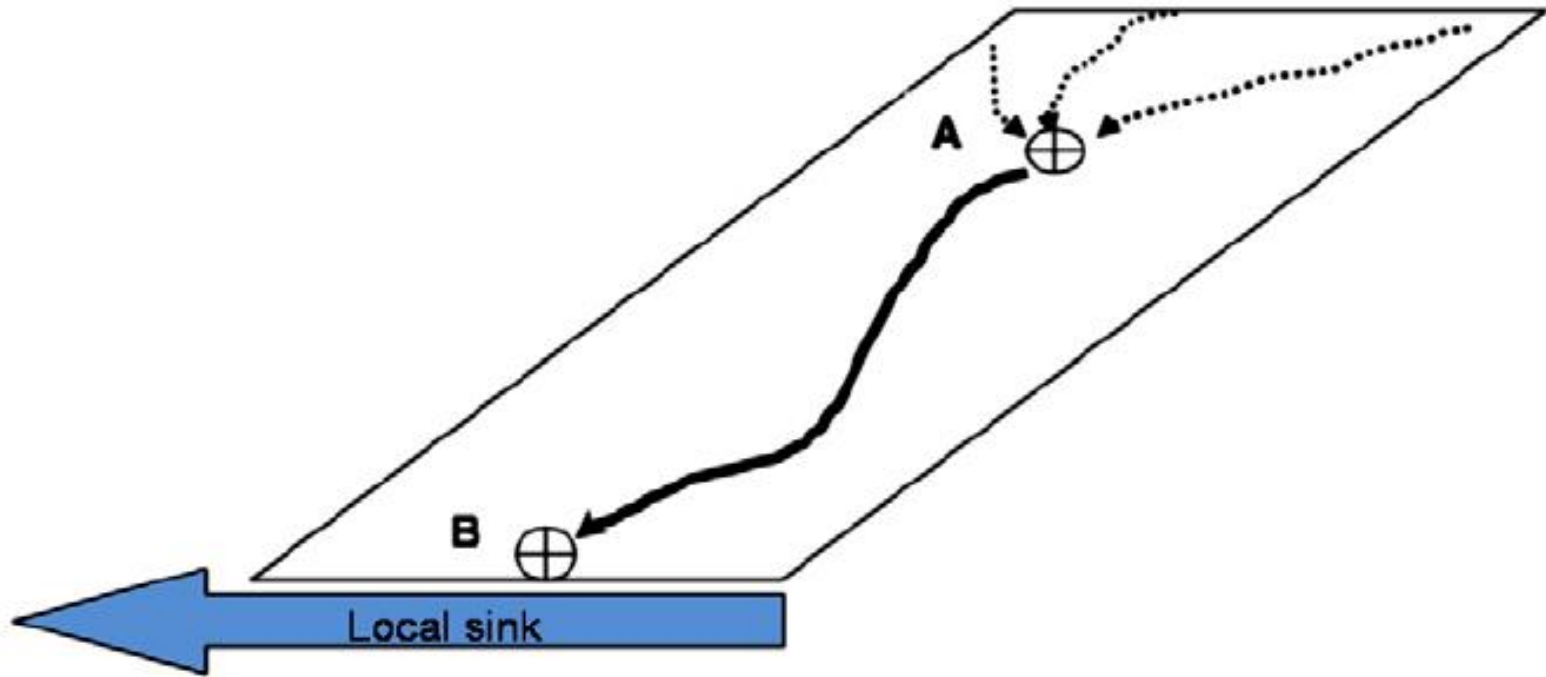
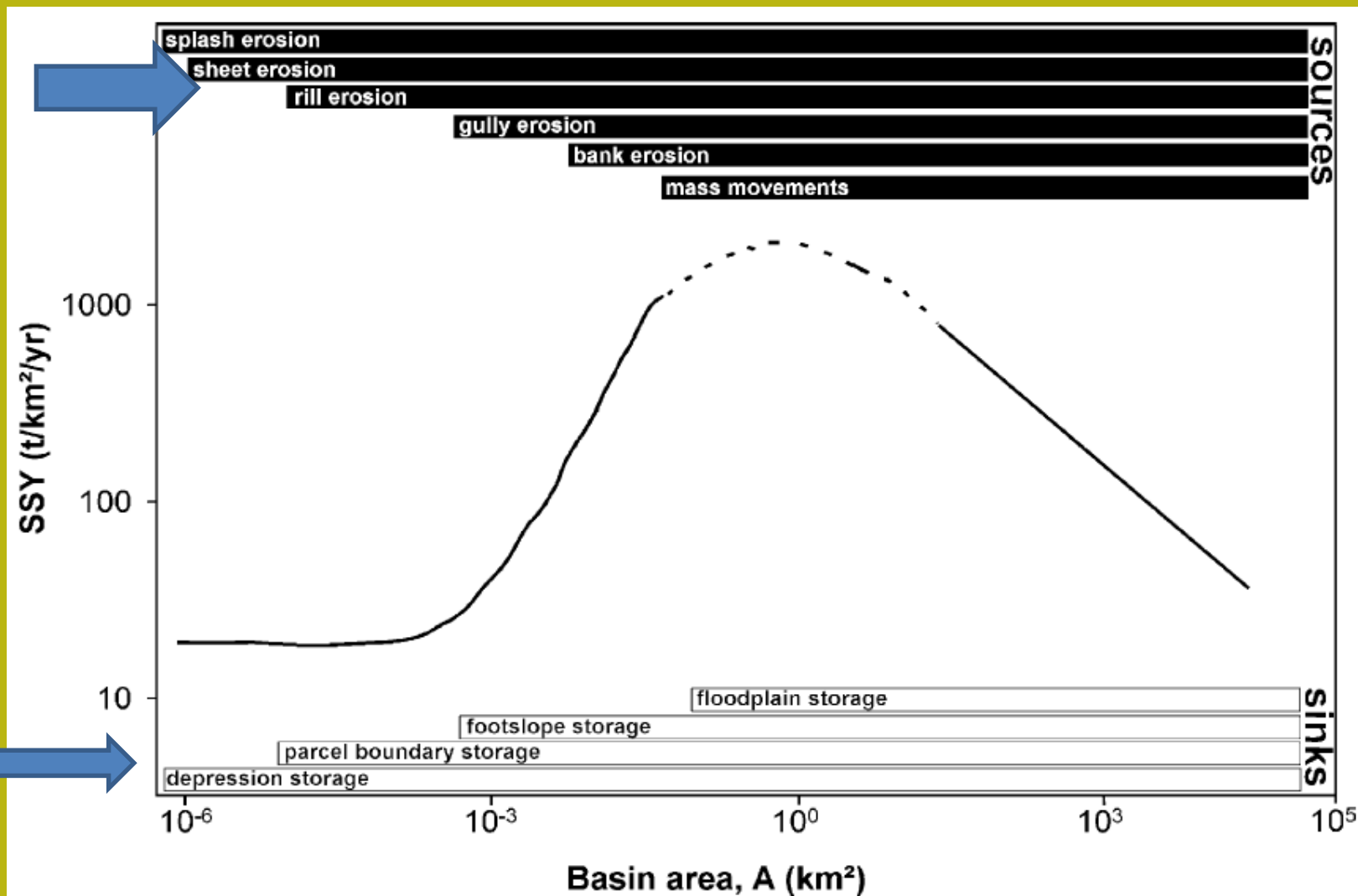


Fig. 1. Simple slope with flow path between the source area A and the local sink B.



# Influence of Connectivity on sediment yield (SSY).....from many sources.... (non only due to soil erosion....)

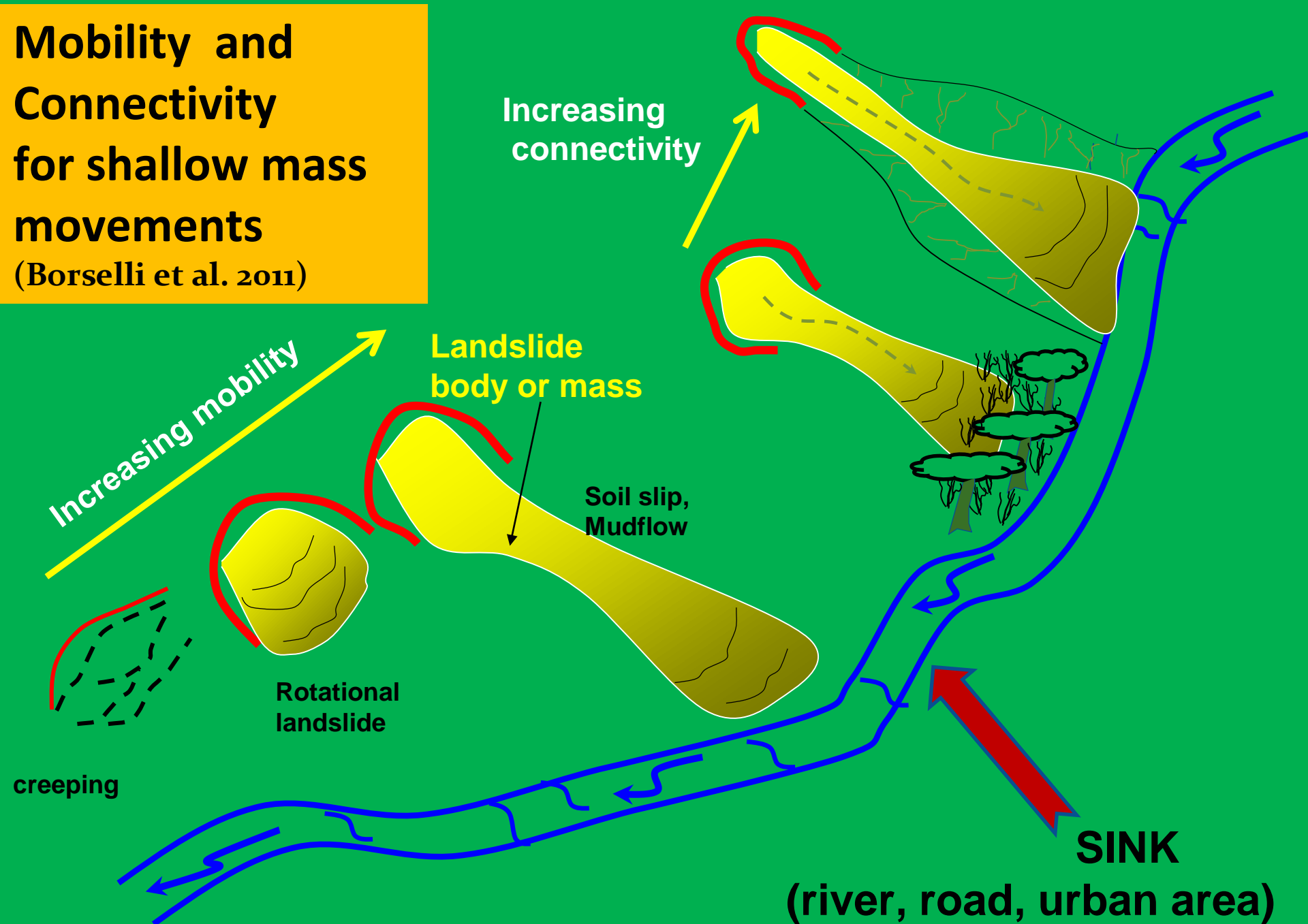
SOURCES



Conceptual model of sediment yield at various scale and contributing sources and sinks (De Vente and Poesen, 2005)

# Mobility and Connectivity for shallow mass movements

(Borselli et al. 2011)

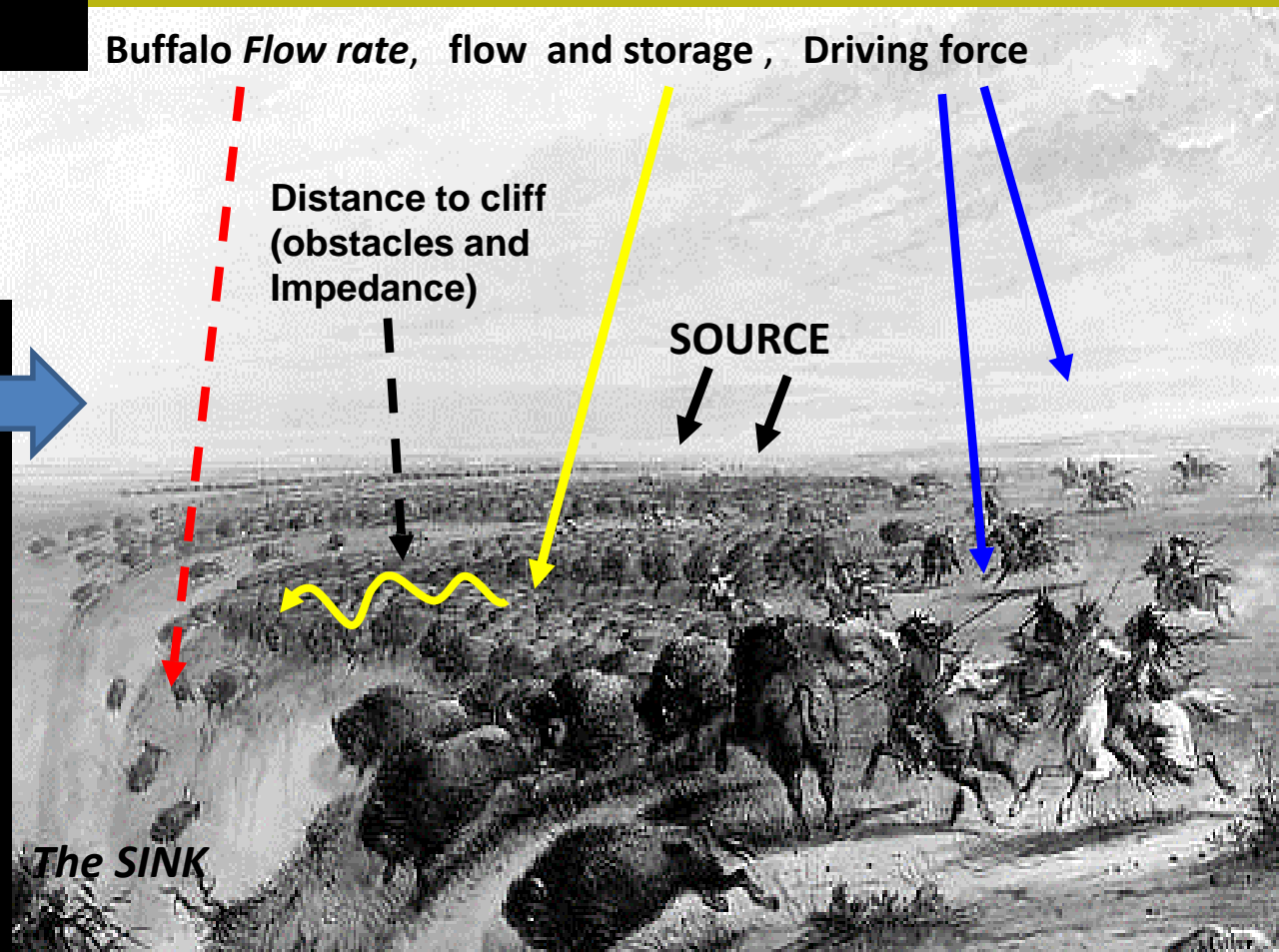




The previous characteristics are defined and used, by many soil erosion distributed models in modelling and computation of erosion and deposition in whole catchments

## A conceptual example... importance of connectivity...

**The Buffalo jump.**  
A Native Americans hunting technique that has some similitude with soil erosion /runoff processes....  
e.g. The chance of each buffalo to fall .... die,  
.. or escape and survive....



Source: Alfred J. Miller 1887 from National Archives of Canada

# Connectivity Indexes, a local metric for flow connectivity approach (FCA).

If an ideal model, able to simulate perfectly mass and energy transfer, redistribution and storage in the landscape, existed, probably we would not need Connectivity metrics.  
*(e.g. A perfect, event based soil erosion model that could simulate erosion, transport and deposition rate on the landscape)*

But this type of models does not exist yet, and many of the existing soil erosion models are not easy to use...sometimes models are extremely complex and are affected by **parametric and modeling uncertainty, numerical problems, and occasionally by numerical instabilities** (e.g. violation of principle of mass and energy conservation).

A set of new tools are needed to consider Connectivity as stand alone metric which can be put in relation to various processes ( e.g. runoff, erosion, mass movement mobility, etc...)

To do this we need to develop Connectivity indexes, as local Metrics, representing the local connectivity status.

But a preliminary and exhaustive *verification in field* of the evidences of the linkage between connectivity and the intensity of certain processes (e.g. soil erosion, soil deposition, landslide mobility, runoff concentration, etc.) is required....





## An attempt to develop a new tool in order to represent a connectivity local metric

### The index of connectivity IC From Borselli et al. 2008

Prolegomena to sediment and flow connectivity in the landscape: A GIS and field numerical assessment

Lorenzo Borselli \*, Paola Cassi, Dino Torri

CNR-IRPI, Via Madonna del Piano 10, 50019 Sesto Fiorentino(FI), Italy

#### In the paper :

Two indices of connectivity were operatively defined:

1) **(IC)** that can be calculated in a GIS environment and represents a map potential connectivity between two different parts of a catchment ( assessment based on landscape's information);

2) Another index that can be evaluated in the field **(FIC)** through direct assessment of connected flow path after a flow or erosion event.

IC and FIC indices were designed to complement each other and their combined use was shown to improve accuracy.

The study was based also on field observation in a 150km<sup>2</sup> watershed Italy and was funded by the European Commission, Directorate- General of Research, **Global Change and Desertification Program, RECONDES project (2004–2007)** “Conditions for Restoration and Mitigation of Desertified Areas using Vegetation” and by **Autorità di Bacino del fiume Arno-Italy; BABI project (2003–2007)**

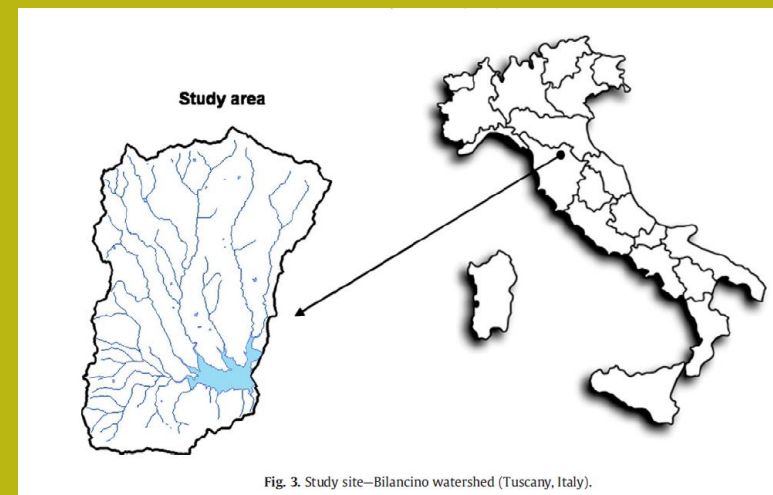


Fig. 3. Study site—Bilancino watershed (Tuscany, Italy).

# IC Index

The Connectivity Index (IC) value is computed using two components:

**Upslope component:** is the potential for down routing due to upslope catchment's areas, mean upslope and land use.

**Downslope component:** is the sinking potential due to the path length, land use and slope along the downslope route.

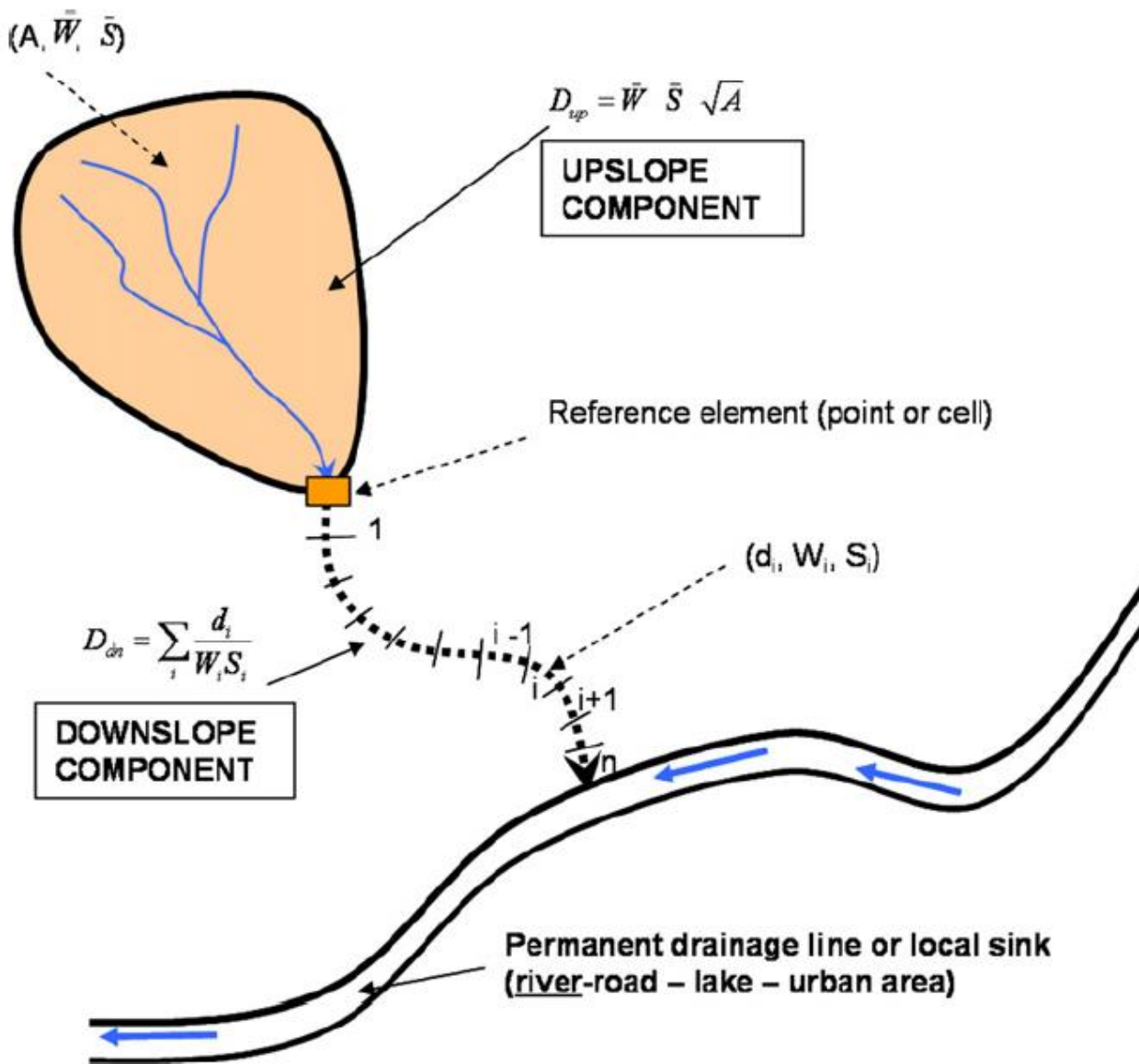


Fig. 2. Definition of IC upslope and downslope component in the landscape for index of connectivity (IC).

$$D_{up} = \bar{W} \bar{S} \sqrt{A}$$



**UPSLOPE  
Component**

$\bar{W}$  = average Weigthing factor in the upslope contributing area (adimensional);

$\bar{S}$  = average slope gradient of the upslope contributing area (m/m)

$A$  = upslope contributing area (m<sup>2</sup>)

$$D_{dn} = \sum_i \frac{d_i}{W_i S_i}$$



**DOWNSLOPE  
Component**

$d_i$  = length of cell  $i$  along the downslope path (in m)

$W_i$  = Weighting factor of cell  $i$  along the downslope path (adimensional)

$S_i$  = slope gradient of cell  $i$  along the downslope path (m/m)

**Final IC calculation in each pixel**



$$IC = \log_{10} \left( \frac{D_{up}}{D_{dn}} \right) = \log_{10} \left( \frac{\bar{W} \bar{S} \sqrt{A}}{\sum_i \frac{d_i}{W_i S_i}} \right)$$



$$IC = \log_{10} \left( \frac{D_{up}}{D_{dn}} \right) = \log_{10} \left( \frac{\bar{W} \bar{S} \sqrt{A}}{\sum_i \frac{d_i}{W_i S_i}} \right)$$

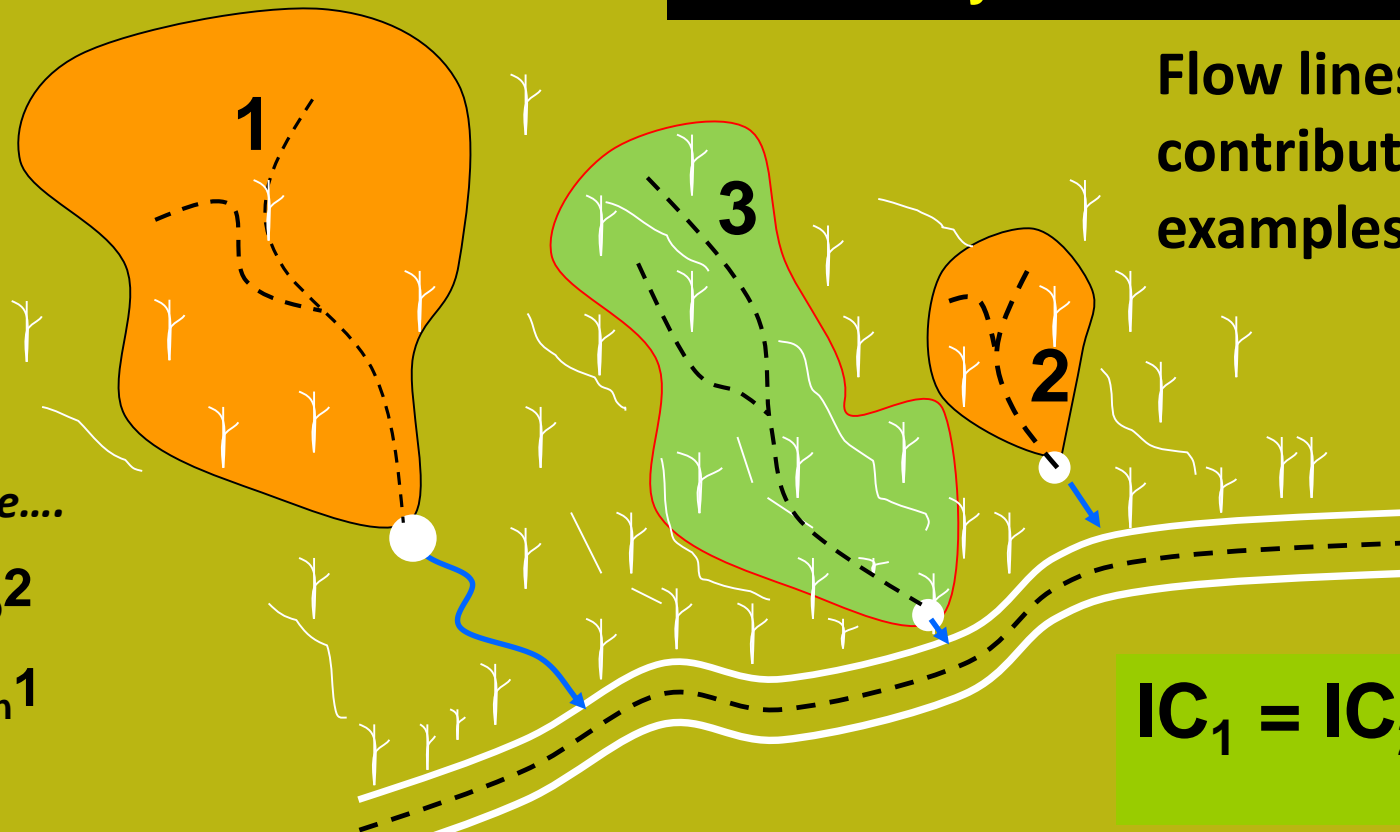
**IC range:**  $[-\infty, +\infty]$

## Orders of magnitude of IC

*Under this definition the local level of connectivity to permanent drainage lines/sinks is inversely proportional to IC:*

Values  $IC > 0$  **high connectivity**

Values  $IC < 0$  **medium to low connectivity**



**Flow lines and contribution area examples**

*A special case....*

$$D_{up}1 > D_{up}2$$

$$D_{dn}2 < D_{dn}1$$

$$IC_1 = IC_2 < IC_3$$

## IC methodology requirements (classical approach):

- High quality and high resolution DTM are preferred. Ideal resolution variable between 2 and 5X5 m
- Resolution of DTM at least 10x10m may be sufficient
- But in some cases may be acceptable until 20X20 depending from local availability and from the type of application we want to generate
- Raster map of slope gradient
- Detailed Land use to obtain the local weighting (W) factor map
- No data value layer (internal local sink): river mask, roads, urban area, lakes, etc..
- River mask must be generated starting from a maximum accumulation area (it defines permanent drainage lines) - usually 1-2.5 ha.

# IC Index computation notes

Permanent drainage lines, roads, urban areas, and water bodies, as well as pixels outside of the considered watershed, are usually set as no data value MASK.

ALL internal *no data value* pixels are considered as local SINKs

The concept of local sink is fundamental...

Of course you can choose to not consider road or urban areas as local SINK, but in this case you will generate an new type of connectivity pattern and values...

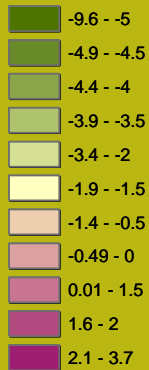
DTM quality and resolution is fundamental... !!!! In any case and at any resolution

## Types of no data value masks

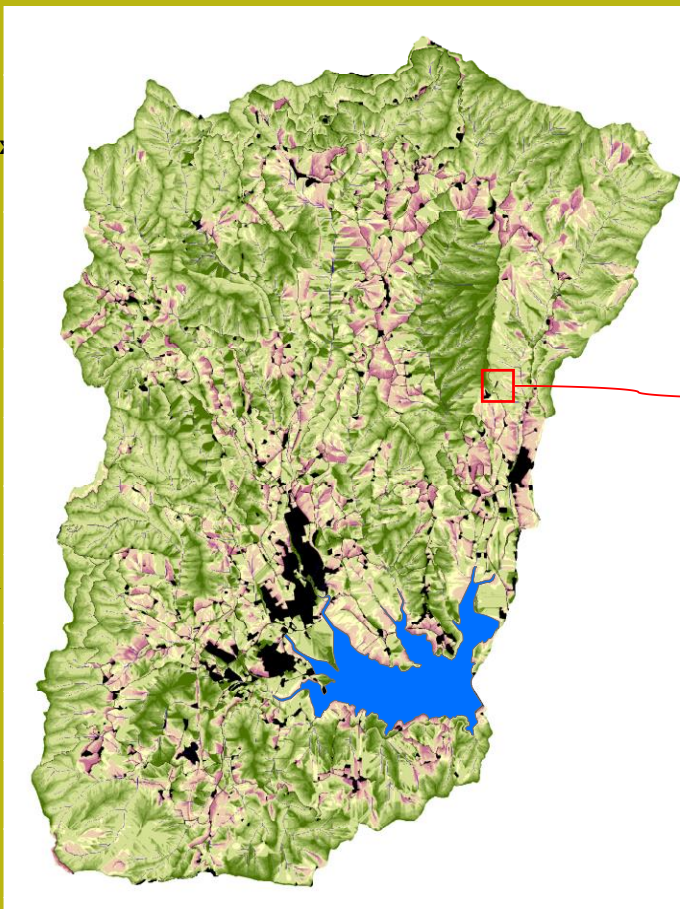




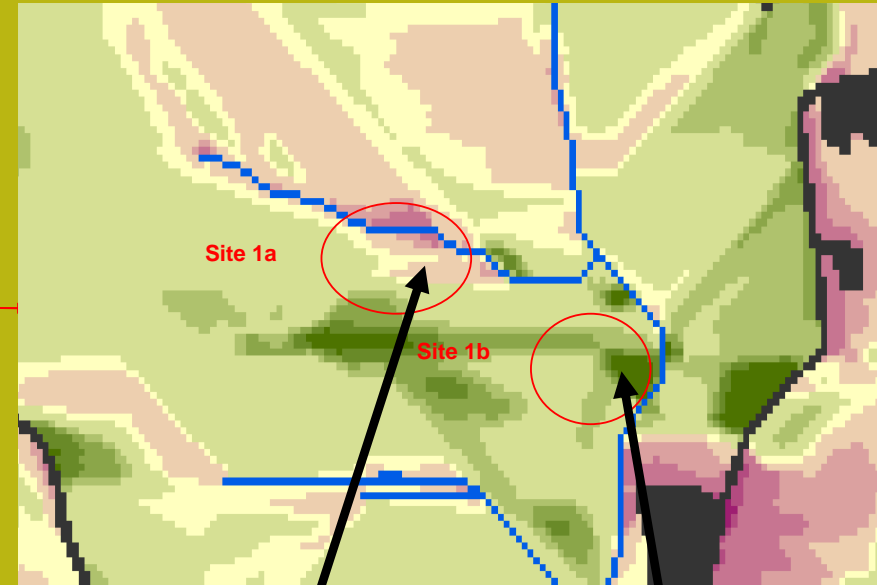
Connectivity fluxes index:



In black the road/urban mask



IC map of deposition and connection areas evidenced in red



Site1a -Area in proximity of local sink at field bottom: direct connection of rill system without appreciable sedimentation



Site1b-Area in proximity of local sink at field bottom: direct connection of rill system relevant sedimentation

**Application at watershed scale  
(Bilancino – Tuscany - Italy):  
hot spot identification of primary  
sediment sources area.**

**From Borselli et. Al. (2007,2008)**

**Area: 150 km<sup>2</sup>**

**DTM 5x5m**

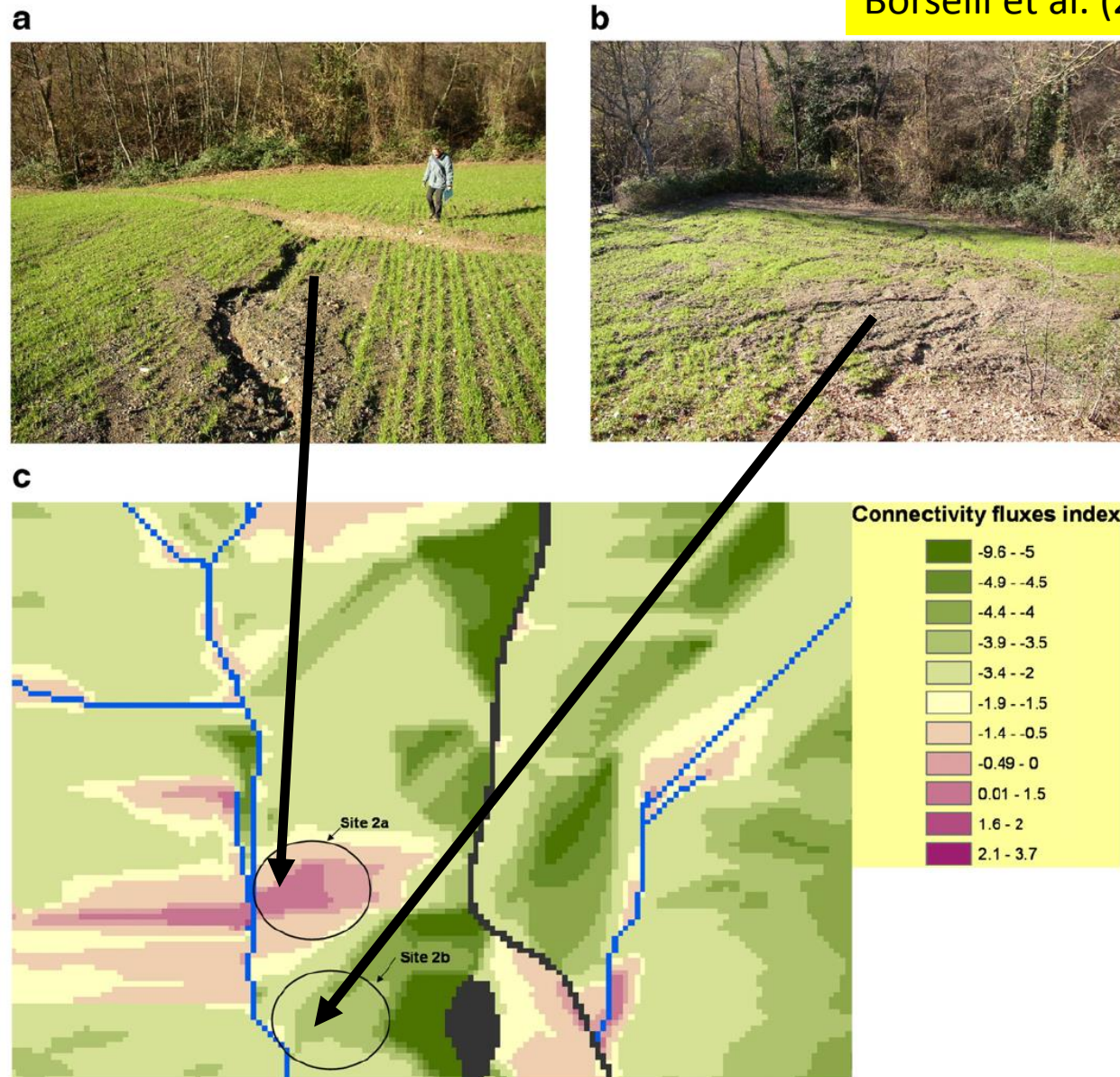


Fig. 5. a: Site 2—Area close to a local sink at the bottom of a field: direct connection of rill system without detectable sedimentation. b: Site 2—Area in proximity of local sink at the bottom of a field: direct connection of rill system with intense sedimentation. c: IC map of Site 2: deposition and connection areas are evidenced inside circular areas.

## FIC – THE FIELD APPROACH

The field assessment of connectivity is evaluated using the same approach of  $I_c$ , i.e. estimating the upslope and the downslope component. A well defined field connectivity index ( $FIC$ ) can permit an evaluation of the GIS-based  $IC$ . As  $FIC$  is strictly related with the event that produced the connectivity evidence, it is also a way of representing ground truth. Hence a comparison between the two indexes will allow for a validation of  $IC$ .

$$FIC = \frac{Su + Sd}{2}$$

$FIC$  is subdivided into an upslope ( $Su$ ) and a downslope subfactor ( $Sd$ );  $FIC$  varies between 1 and 100 and increases with connectivity

$$Su = Au + Bu + W_u Cu_1 + (1 - W_u) Cu_2$$

Upslope subfactor

$$Sd = Ad + Bd + Cd + W_{d1} Dd_1 + (1 - W_d) Dd_2$$

Downslope subfactor

**Au** and **Ad**- Upslope area and the downslope distance to local sink;

**Bu** and **Bd** - Presence of sedimentation features along the upslope and downslope flow paths;

**Cd** presence and wideness of buffer-bush strip area before or along the flow path to the local sink;

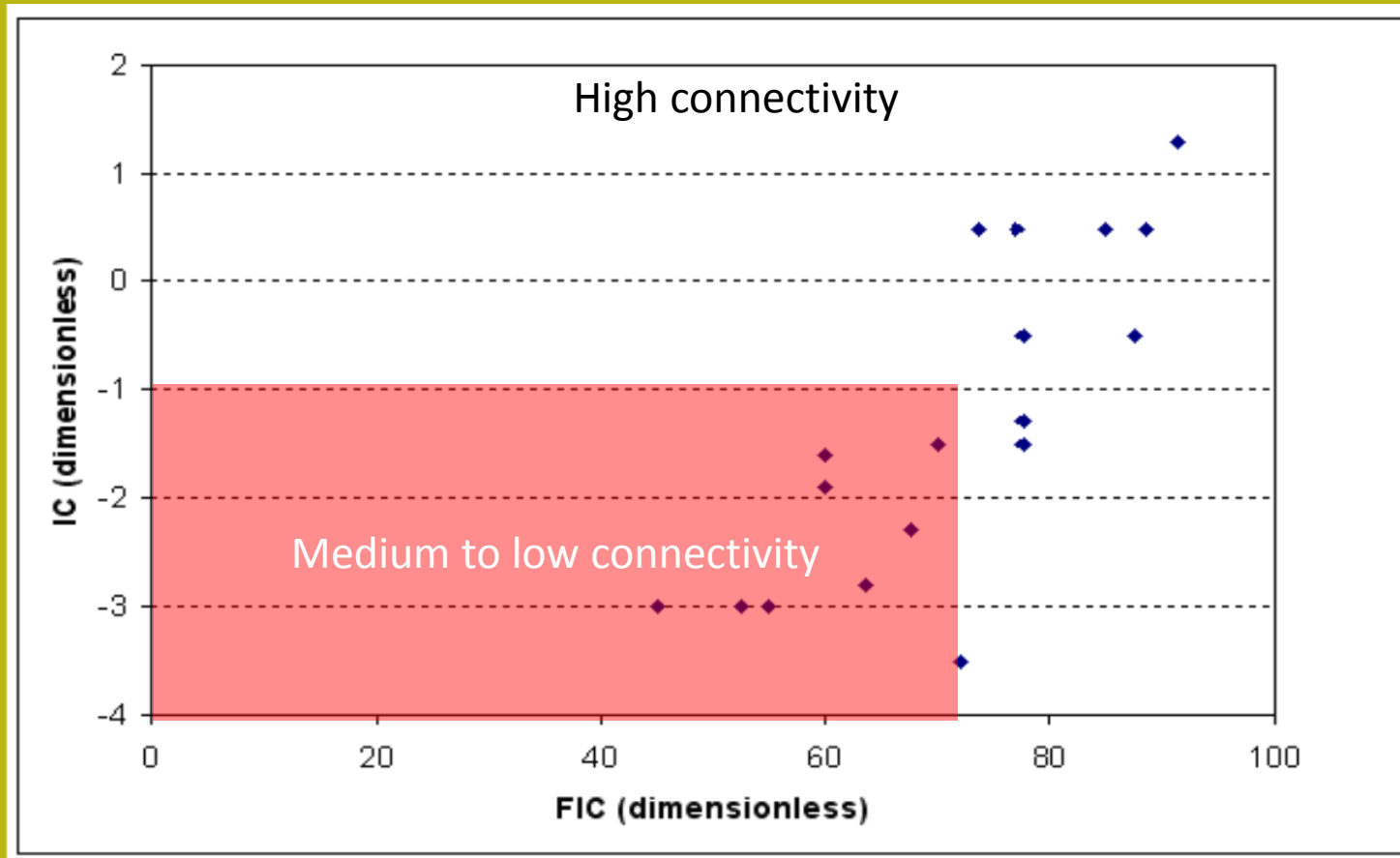
**Cu1**, **Cu2**, **Dd1**, **Dd2** - Subfactors opposing resistance to fluxes (surface roughness and vegetation cover crop, plant basal area) in the upslope area and the downslope flow path;

**Wu** and **Wd** – Fraction of arable land in the upslope area and in downslope flow path.

From Borselli et al.(2008)

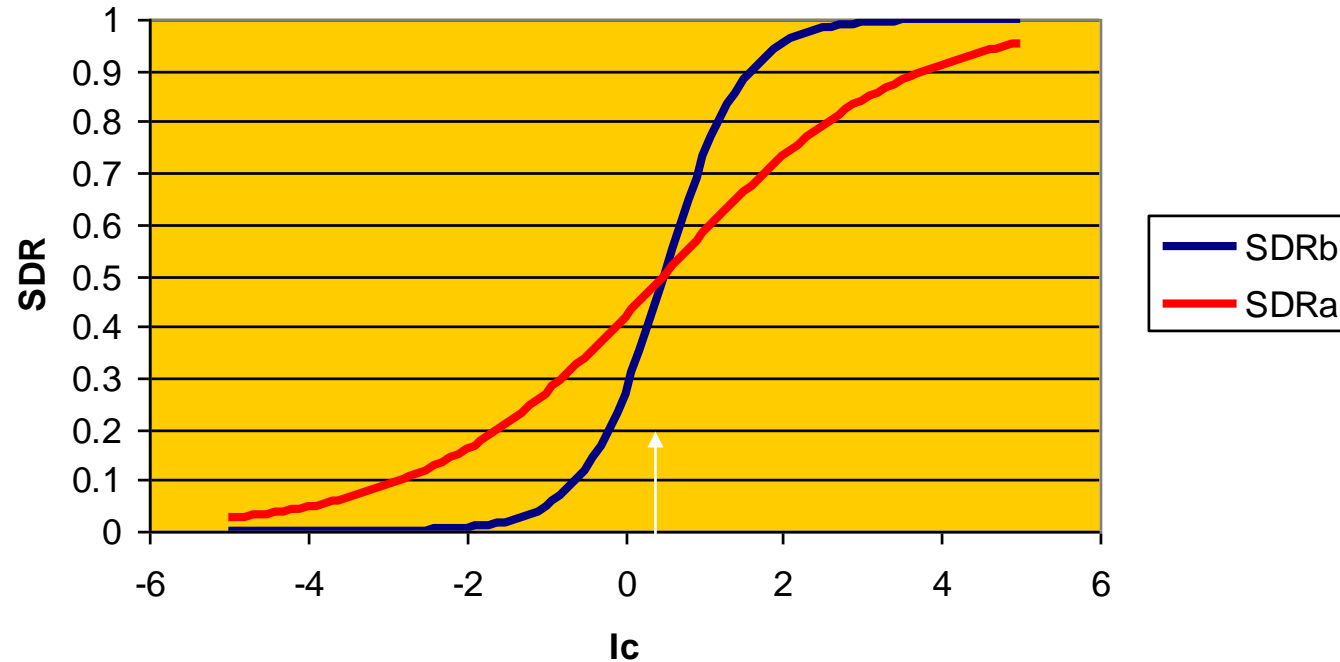
**Please see the original paper for details and examples of application...**





The FIC values have been compared to the IC flux map obtained with the ArcGIS procedure for the entire study site

## Possible functional relationships between IC and SDR...



**Boltzmann type function:  $SDR=f(IC, IC_0, k)$**   
(Borselli et al. 2007)

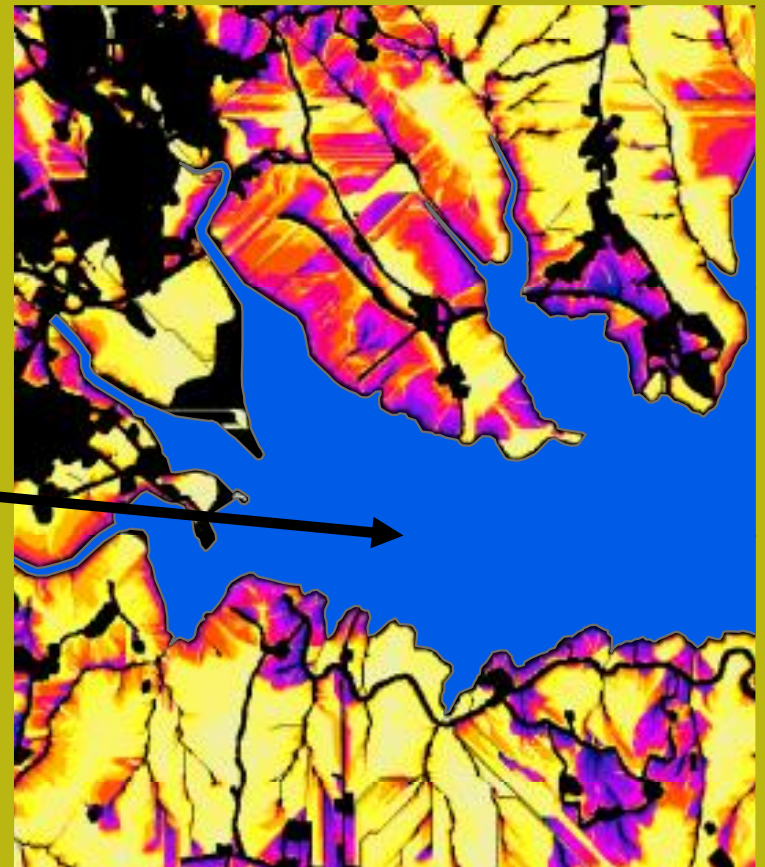
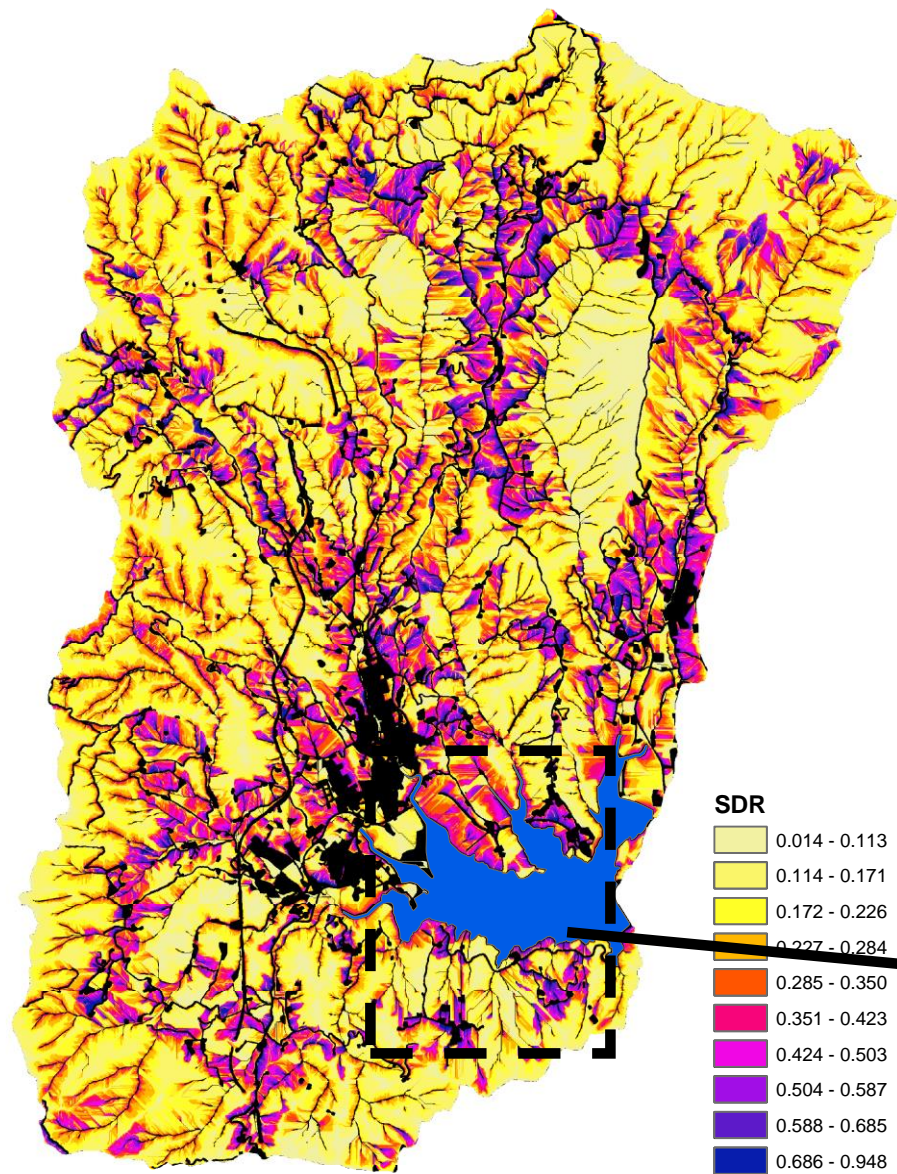
$$SDR = \frac{1}{1 + \exp\left(\frac{IC_0 - IC}{k}\right)}$$

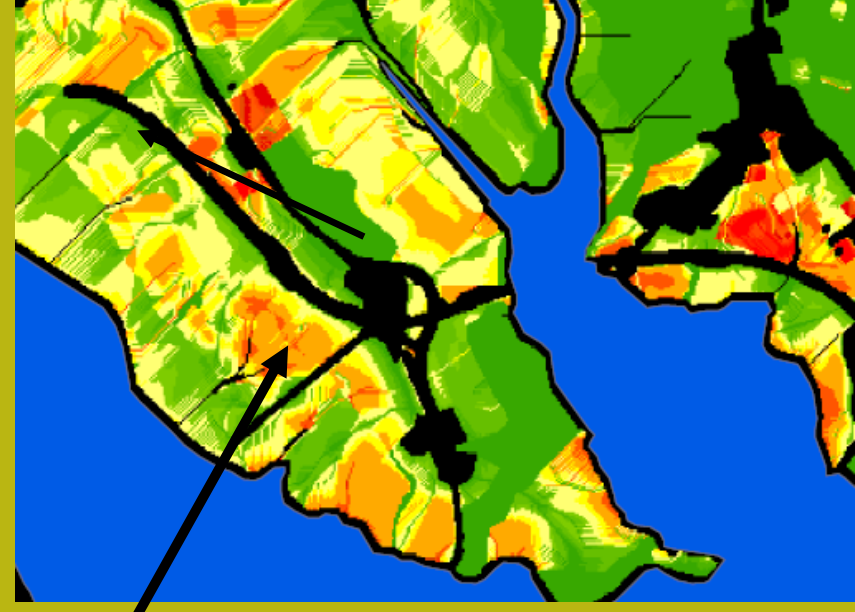
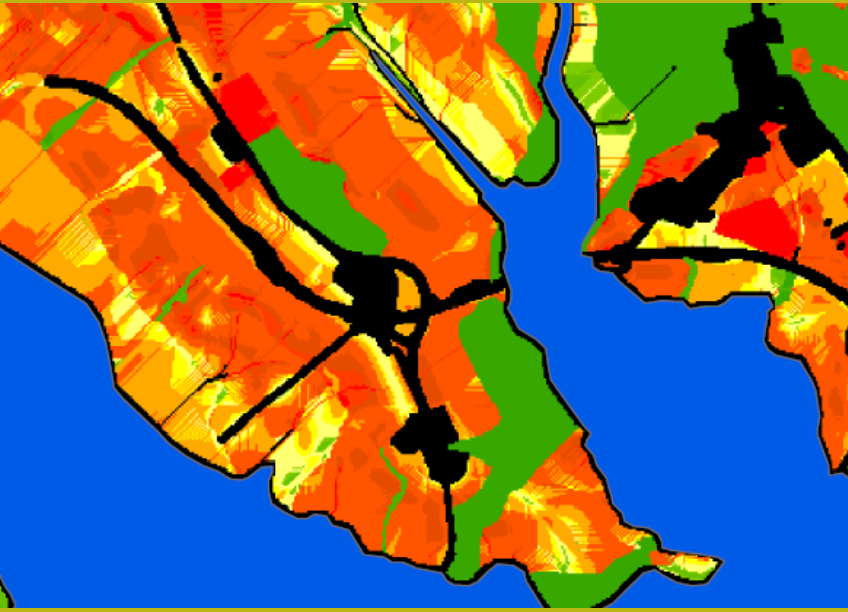
In this first application (2007) to Bilancino watershed we speculated on a possible relationship between IC and SEDIMENT DELIVERY RATIO (SDR) and developed it as  $SDR=f(IC)$  at the hillslope scale.

This analysis was not officially published until 2012 after the verification made by Vigiak et al. (2012) in another catchment (in Australia).

## SDR map in a watershed for correction of average annual erosion rate (under USLE-type models )

Many authors have used the SDR to correct for distributed soil erosion model outputs (Ferro and Porto 2000; Lu et al., 2006)





## Local average erosion rate... Classic RUSLE 2D

In the Bilancino application  
a new algorithm for the calculation  
of erodibility (k) based on global dataset and  
climatic classification was also used ...

KUERY software 1.2.

See Borselli al. (2009,2012)

[www.Lorenzo-borselli.eu/kuery](http://www.Lorenzo-borselli.eu/kuery)

Average sediment yield  
contribution:  
RUSLE2D corrected  
according to IC and SDR

**Primary sources of sediment  
are in red (higher erosion rate)**



# Proposals for extensive applications of IC (....and FIC) *in 2008*:

The proposed procedure for the IC model contained a large set of potential applications such as:

- 1) hot spot identification of primary sediment sources to permanent drainage lines;
- 2) verification of effects of eco-compatible mitigation measures to reduce or favor connectivity (Hooke and Sandercock, 2007);
- 3) monitoring changes in the degree of connectivity in areas with high geomorphological evolution rates;
- 4) Performing scenario analysis to assess efficiency of conservation measures against soil erosion, sediment, and nutrients transport, and siltation (all strongly related to flux connectivity).

## A first Tool :

In the paper there is a sequence of commands for arcMap (ArcGIS 8.3) was provided to facilitate the calculation of IC

The procedure is not yet outdated....

even if now exists some other valid and more rapid alternatives (e.g. Connectivity Toolbox , Cavalli et al. 2014 see forward..)

***Pseudocode fragment for IC calculation - Borselli et al. 2008***

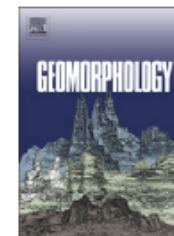
### Appendix A. ArcGIS 8.3, ArcMap (spatial analyst extension)

Given data  
grid: elevation, C  
shapefile: road, urban area  
**Computation of input data**

1. Slope without null value
  - a. Enable Spatial Analyst  
under View... Toolbars, select Spatial Analyst
  - b. Calculate Slope  
from the Spatial Analyst toolbar, select Surface Analysis... Slope  
name the new theme Slope
  - c. Raster Calculator  
from the Spatial Analyst toolbar, select Raster Calculator  
build an expression  $(([Slope]=0) * 0.005) + [Slope]$ ;  
name the new theme S
2. Road/urban mask
  - a. create a raster map of road and urban areas  
from the Spatial Analyst toolbar, select Convert.. feature to raster  
....

**Other applications of connectivity index IC were found in the international scientific literature, after 2008 .... Mainly after 2010....**

**Only publications in ISI journals area have been considered in following slides....**



## Comparison of conceptual landscape metrics to define hillslope-scale sediment delivery ratio

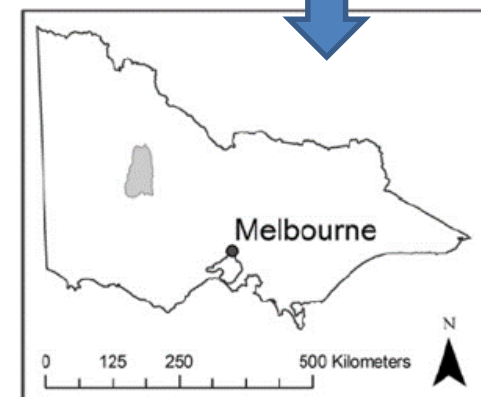
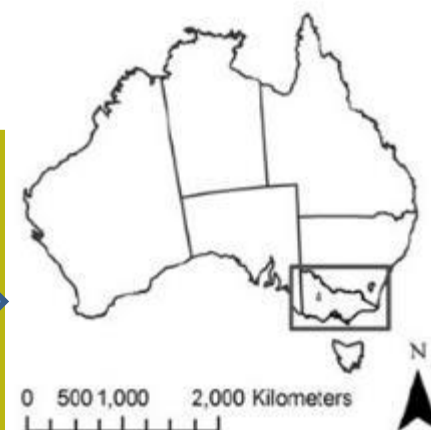
O. Vigiak <sup>a,\*</sup>, L. Borselli <sup>b</sup>, L.T.H. Newham <sup>c</sup>, J. McInnes <sup>a</sup>, A.M. Roberts <sup>a</sup>

**Site :** Avon-Richardson catchment (Victoria, Australia) 3300 km<sup>2</sup> .. **DEM:** Raster DEM 20x20 m

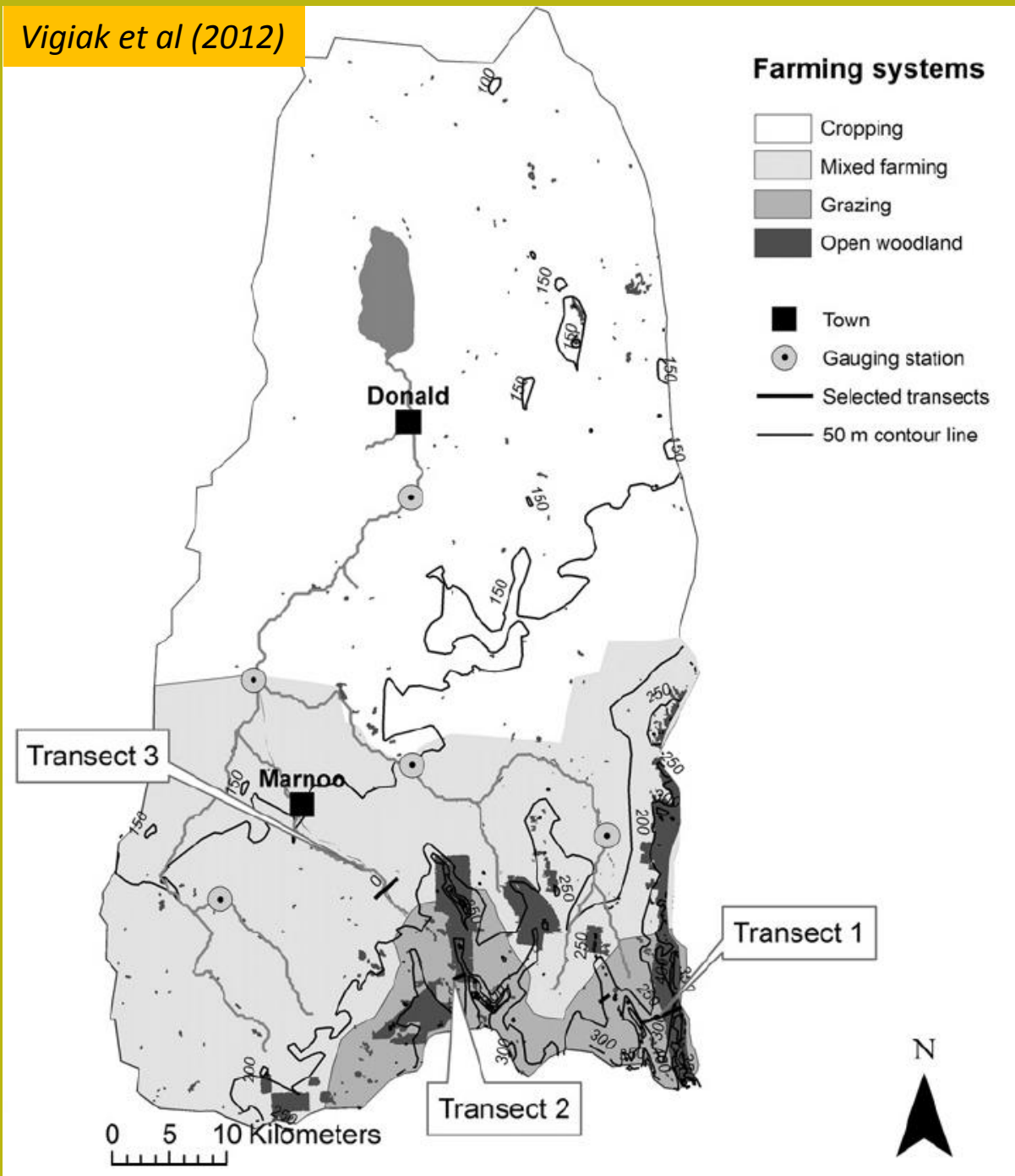
### AIMS:

To calibrate and apply a point-to-catchment linked model (**HowLeaky + CatchMODS**) for daily soil loss estimation (developed by Olga Vigiak and collaborators)

**4 metrics for regionalization of SDR were compared in terms of pattern distributions and efficiency in matching sediment yields at 5 monitoring stations (4 indices of connectivity in total, including IC)**




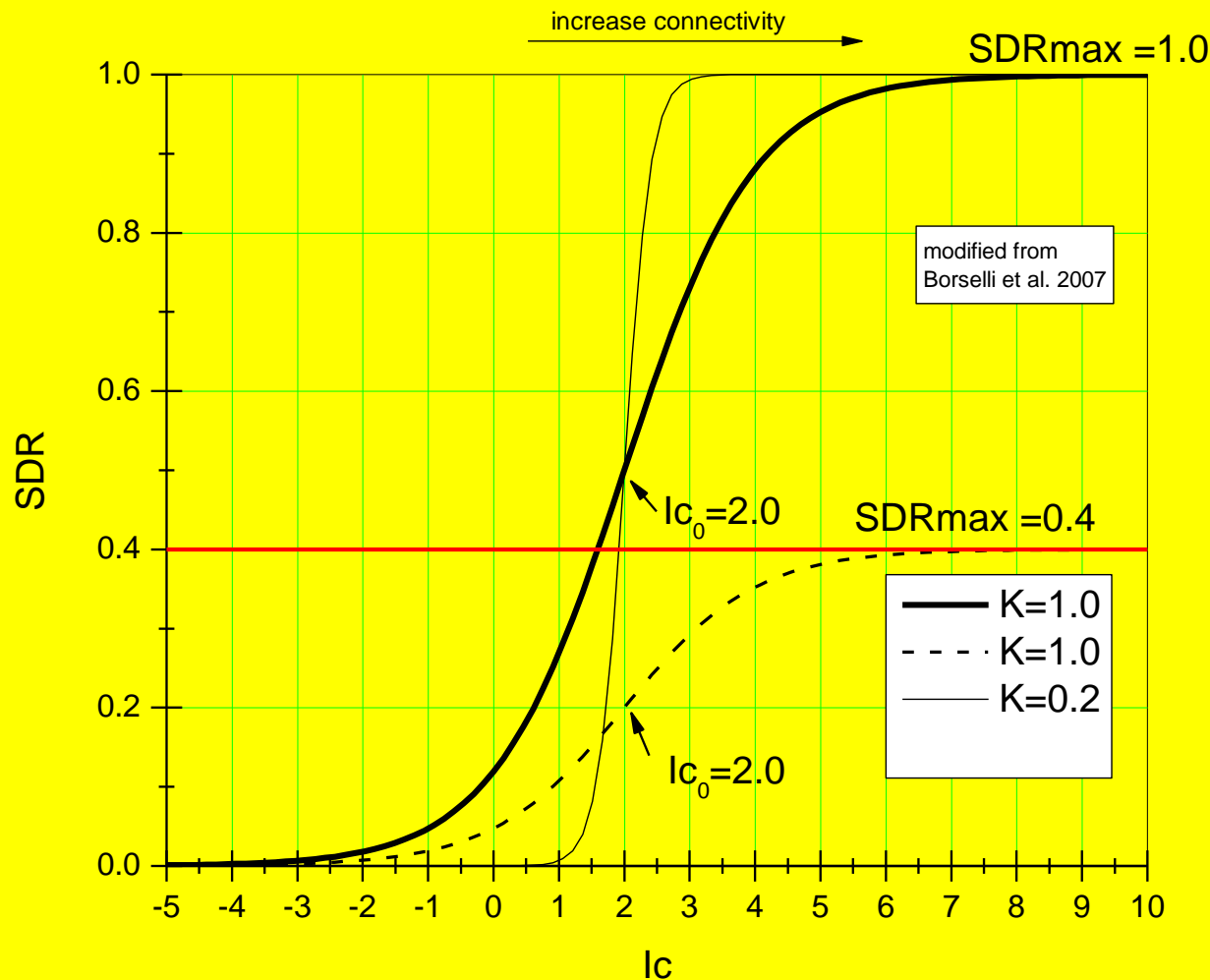
- Avon-Richardson catchment (Victoria, Australia)





# Comparison of 4 connectivity metrics...

Metric	Main concepts
Travel Time (TT, Ferro and Minacapilli, 1995)	(time) distance to the stream
Stream Transport (ST, based on Rustomjii and Prosser, 2001)	Stream transport capacity (Upstream accumulation area and local slope)
<b>Flux Connectivity Index (IC, Borselli et al., 2008) linked to SDR</b> 	<b>Potential for down routing of runoff vs potential for sinks to the stream</b>
Sediment Residence Time (RT, Lu et al., 2006)	Travel time vs effective storm duration, sediment settling properties



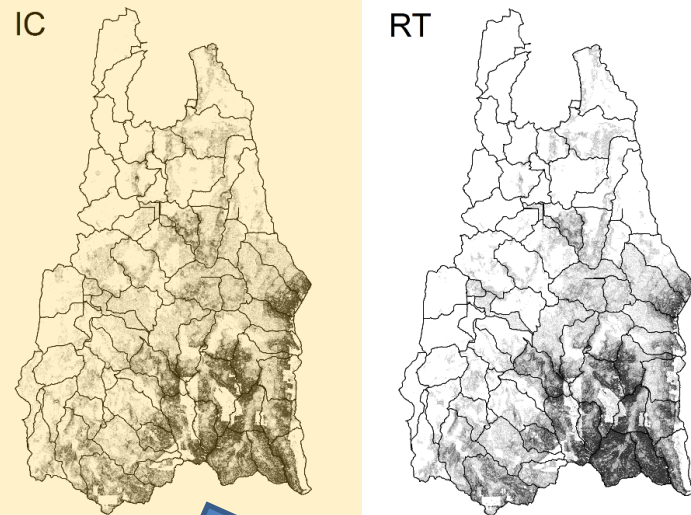
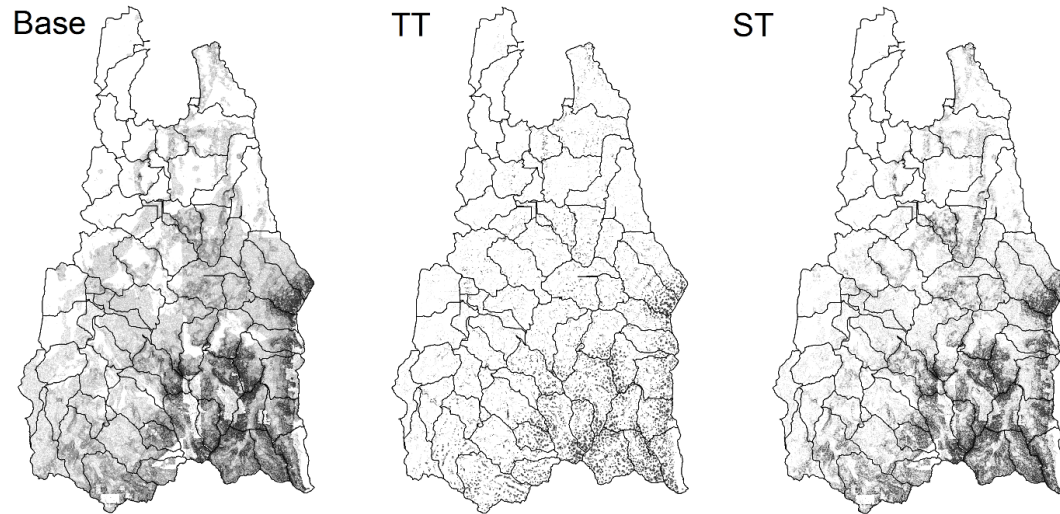
**Implementation  
of  $SDR = f(\text{conn. Index})$   
For 4 types of  
connectivity  
indexes  
(including IC...)  
and optimization  
of parameters  
( $Ic_0$  and  $k$ ) by  
Observed SSY at  
a series of  
Gauging stations**

$$SDR = \frac{SDR_{Max}}{1 + \exp\left(\frac{Ic_0 - Ic}{k}\right)}$$



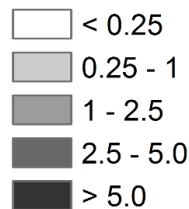
**Case of SDR by IC, and optimized  $Ic_0$  and  $k$   
 $Ic_0=0.5$  and  $k=2.0$**

# Hillslope erosion patterns (pixel)



Specific sediment yield

Mg/km<sup>2</sup>/y



Base = spatially-constant  
SDR

TT = travel time  
(Ferro and Minacapilli  
1995)

ST = sediment transport  
(Rustomjii and Prosser,  
2001)

IC = flux connectivity index  
(Borselli et al. 2008)

RT = residence time  
(Lu et al 2006)

## Estimation of specific sediment yields

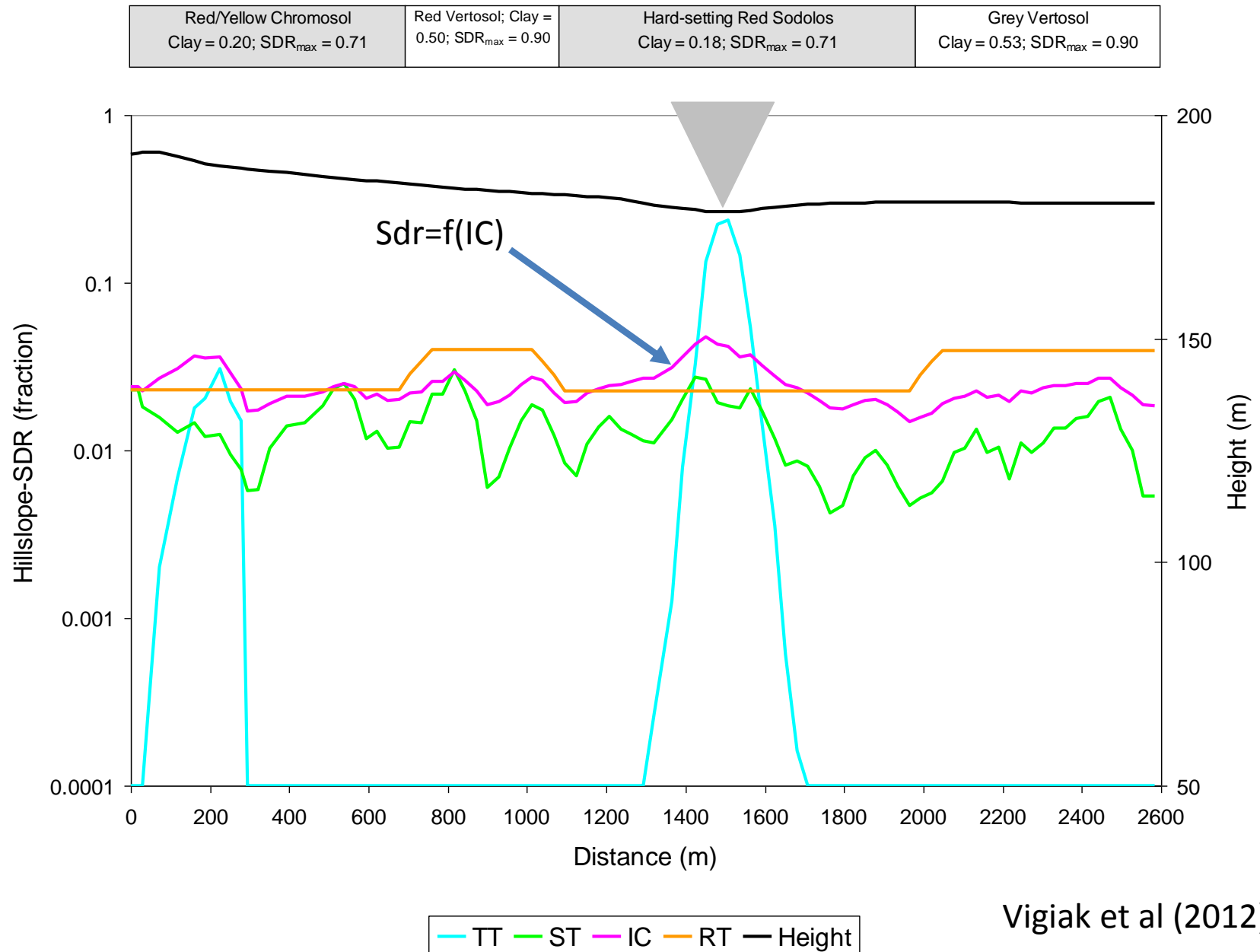
Goodness of fit of  
best model  
configurations at  
subcatchment scale  
against  
measurements at  
**monitoring gauging  
stations (sample  
size=7)**

	Nash-Sutcliffe Model Efficiency	AICc
Base	0.48	-52
TT	-0.25	-46
ST	0.55	-50
IC	<b>0.66</b>	<b>-52</b>
RT	0.61	-54

IC

By Vigiak et al (2012)

# SDR pattern along a transect



Vigiak et al (2012)



# Conclusions (Vigiak et al. 2012)

Regionalisation of hillslope SDR improved the estimation of specific sediment yields at subcatchment scale (less so at pixel scale)

The introduction of all metrics (except RT) did not increase data requirements

The 4 metrics differ in data requirements, dominance of landscape factors, and conceptualization of sedimentological connectivity

**IC metric can be recommended in small-medium catchments (homogeneous climatic conditions)**

**RT metric can be recommended on large catchments (e.g. continental scale; important climatic gradient)**



## Review

# Low erosion rates measured for steep, sparsely vegetated catchments in southeast Spain

N. Sougnez \*, B. van Wesemael, V. Vanacker

George Lemaître Centre for Earth and Climate Research (TECLIM), Earth and Life Institute (ELI), Université Catholique de Louvain, B-1348 Louvain-la-Neuve, B

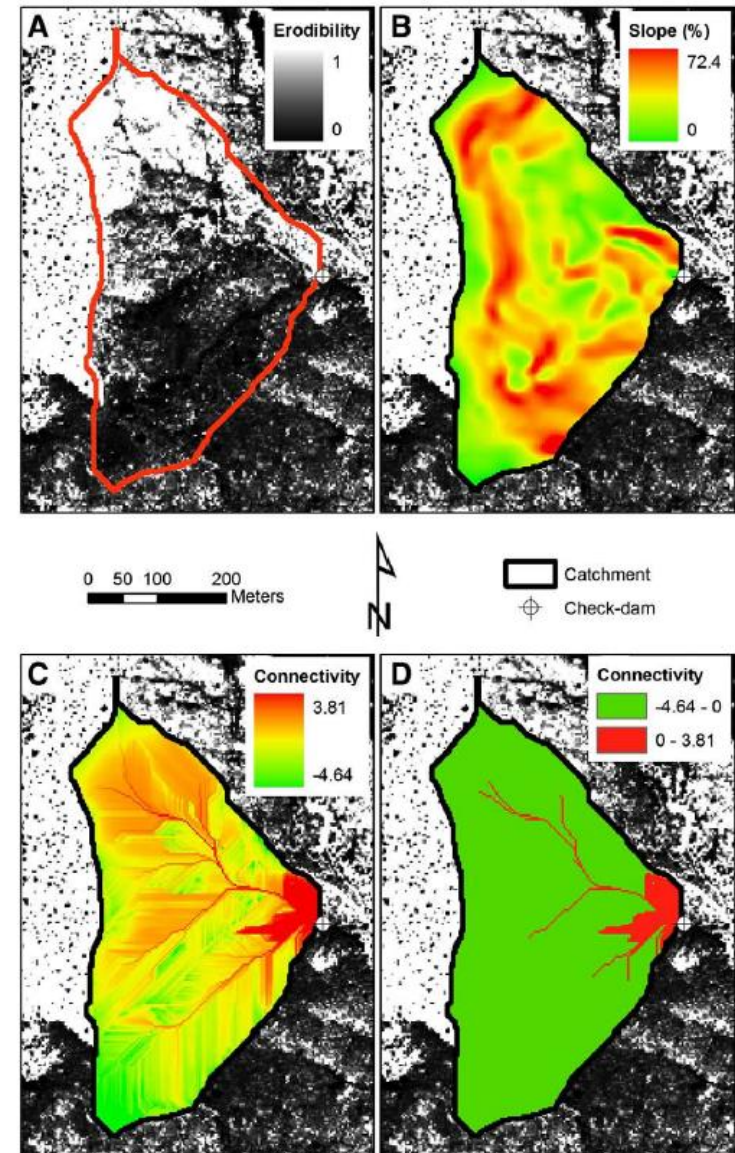
**Site:** Sierra de la Torrecilla and the Sierra de Carrascoy (Murcia, South Spain) – several small catchments of 1–20 ha each.

**DTM:** 3x3 m from contour lines, topographic Map 1:25.000

**aims:** to provide an **accurate estimation of catchment-wide erosion rates for a semi-arid mountainous region**. A variety of methods combined to measure and analyze spatial **patterns in vegetation cover; and to evaluate their effect on water and sediment connectivity**



N. Sougnez et al. / Catena 84 (2011) 1–11





ELSEVIER

Contents lists available at SciVerse ScienceDirect

Catena

journal homepage: [www.elsevier.com/locate/catena](http://www.elsevier.com/locate/catena)

Year 2013

## Predicting runoff and sediment connectivity and soil erosion by water for different land use scenarios in the Spanish Pre-Pyrenees

M. López-Vicente <sup>a,\*</sup>, J. Poesen <sup>a</sup>, A. Navas <sup>b</sup>, L. Gaspar <sup>b</sup>

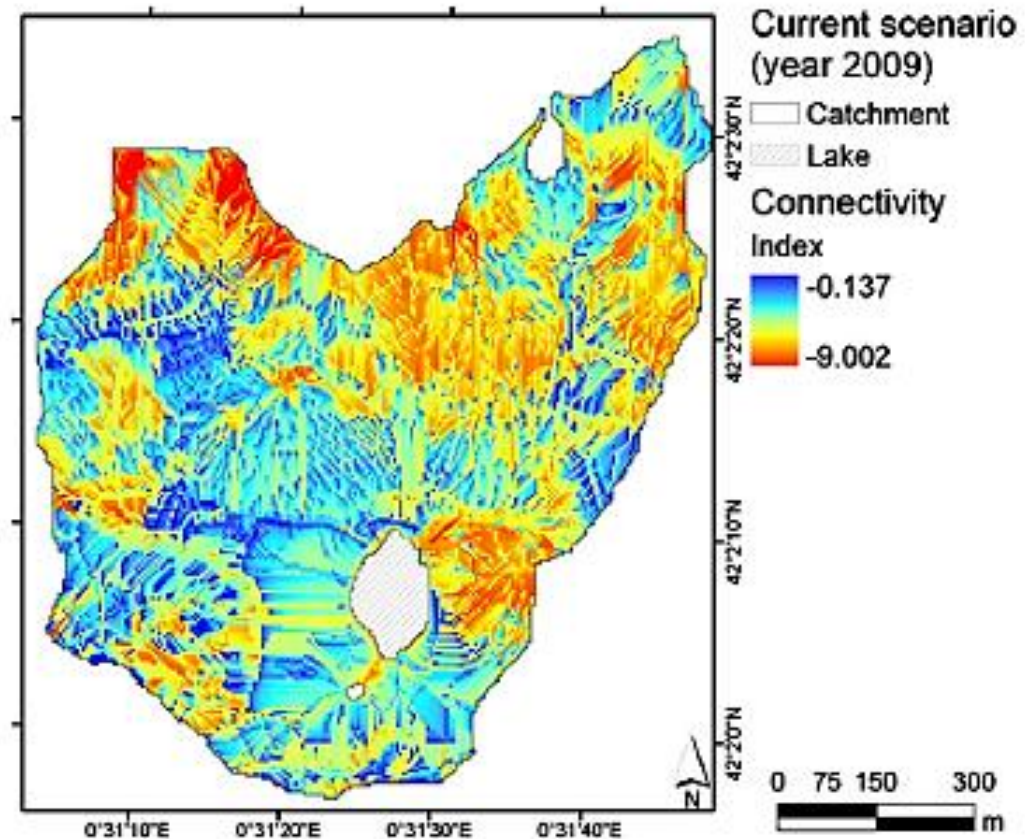
<sup>a</sup> Dept. of Earth and Environmental Science, K.U. Leuven, GEO-INSTITUTE, Celestijnenlaan 200 E, 3001, Heverlee, Belgium

<sup>b</sup> Dept. of Soil and Water, Experimental Station of Aula Dei, CSIC, Postal Box 202, 50080, Zaragoza, Spain

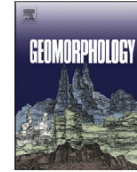
**Site:** Spanish Central Pre-Pyrenees.  
0.74 km<sup>2</sup> catchment

**DTM:** not provided (probably 5x5m)

**aims:** the study seeks to assess the effect of agricultural terraces, irrigation channels, trails, sinks, scarps, and land abandonment on the hydrological connectivity of a small catchment and its consequences on predicting rates of soil erosion under four different scenarios of land uses.





**Lithology:** mica-schist, gneiss, and phyllite.**Land use:** coniferous forest, mountain grassland, bare rock and debris.**Annual rainfall:** around 500 mm in the valley floor, strong increase with elevation.**Gadria catchment :** area 6.36 km<sup>2</sup>, average slope 79.1%, range in elevation: 1394 – 2945 m.**Strimm catchment :** area 8.5 km<sup>2</sup>, average slope 61.8%, range in elevation 1394 – 3197 m.

## Geomorphometric assessment of spatial sediment connectivity in small Alpine catchments

Marco Cavalli <sup>a,\*</sup>, Sebastiano Trevisani <sup>b</sup>, Francesco Comiti <sup>c</sup>, Lorenzo Marchi <sup>a</sup>

<sup>a</sup> CNR-IRPI, National Research Council of Italy - Research Institute for Geo-Hydrological Protection, Padova, Italy

<sup>b</sup> University IUAV of Venice, Faculty of Architecture, Venezia, Italy

<sup>c</sup> Free University of Bozen-Bolzano, Faculty of Science and Technology, Bolzano, Italy

**Sites:** Gadria and Strimm catchments  
(Eastern Italian Alps) - 14.4 km<sup>2</sup>

**DTM:** 2.5x2.5 m (high resolution)

**aims:** development and **adaption of IC index to model sediment pathways dealing with debris flows and channelized sediment transport**, based on the one proposed by Borselli et al. (2008) with *ad hoc* modifications aimed at better exploitation of HR-DTMs

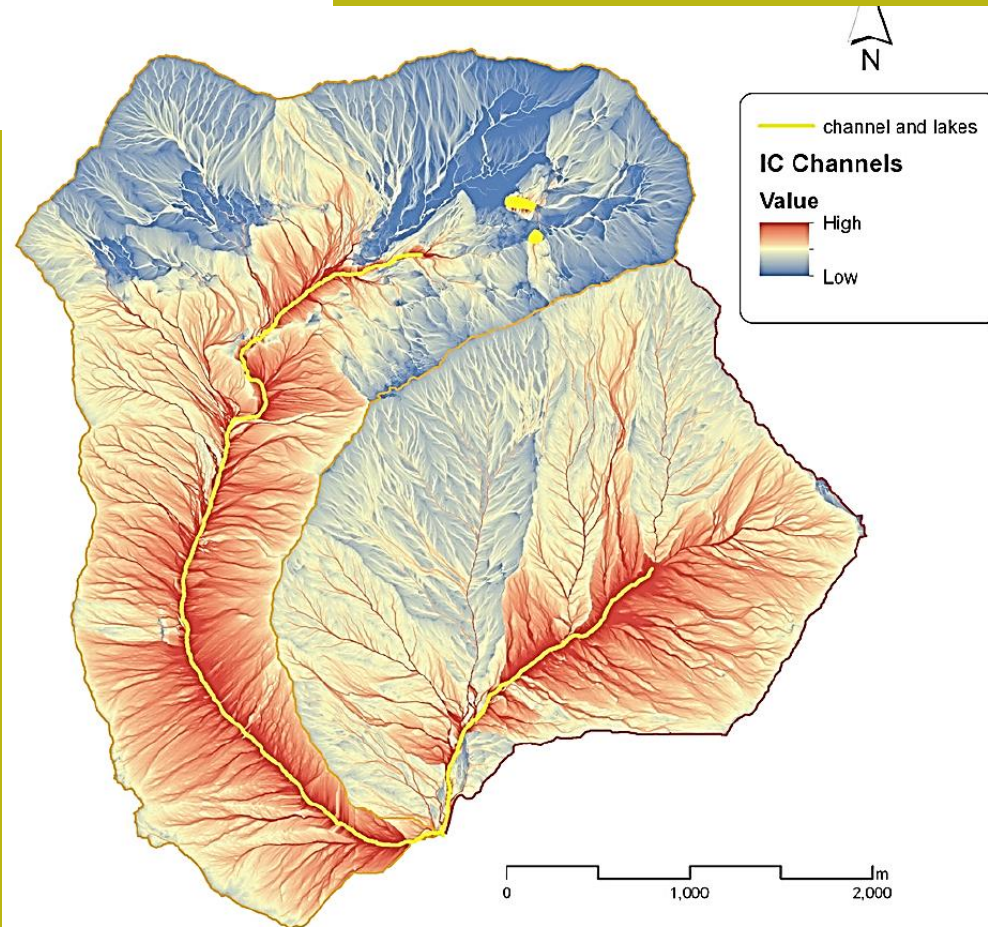
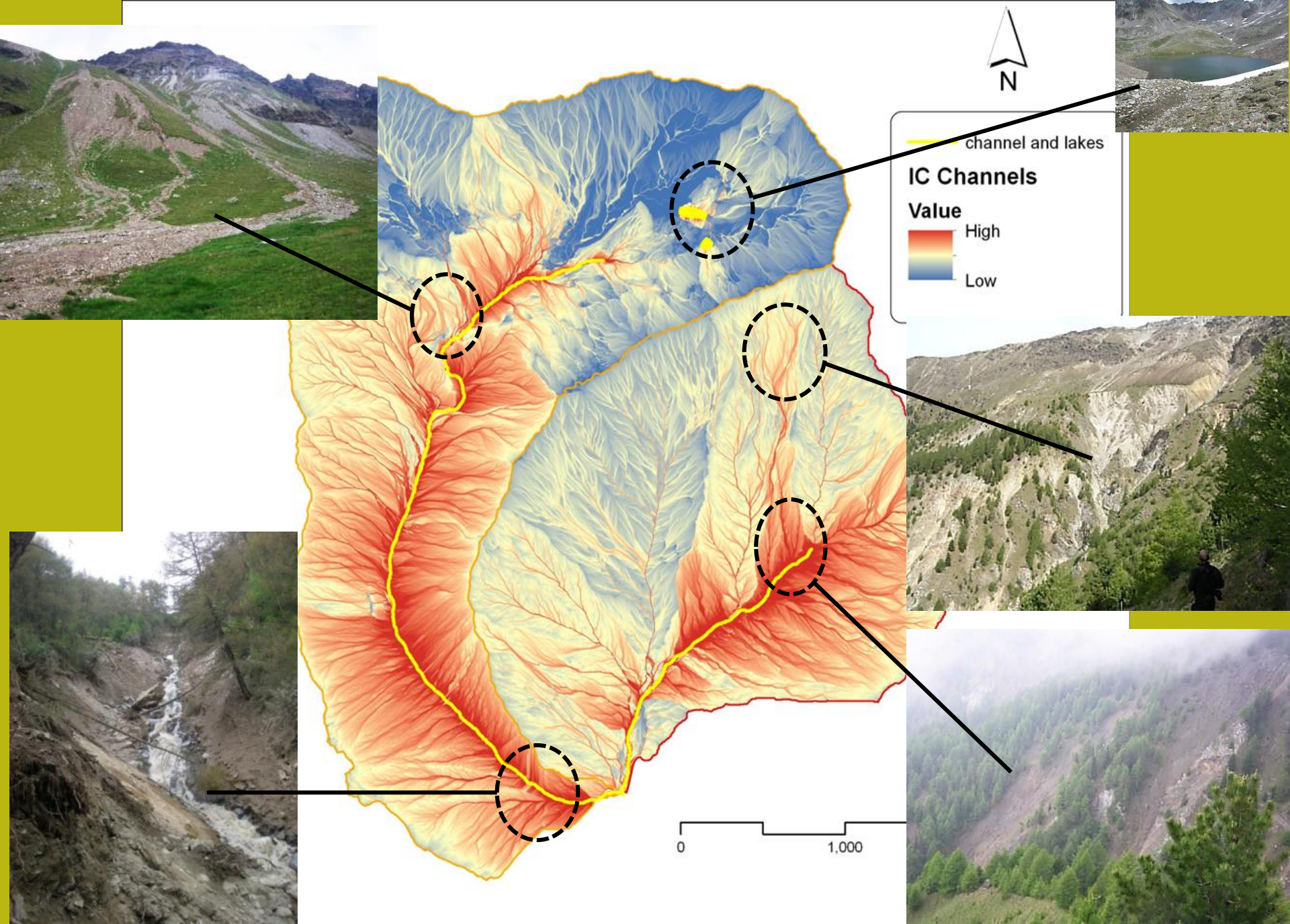


Fig. 5. IC channels map: index of connectivity IC computed with reference to main channels and lakes.







# Adaptation of the IC to mountain catchments and its use with HR-DTMs (Cavalli et al., 2013)

**BEWARE !!!**  
High resolution DTM is required....

## Flow direction

$D \infty$  method (Tarboton, 1997)

## Slope $S$ (m/m)

$S > 1 \quad \rightarrow \quad S = 1$

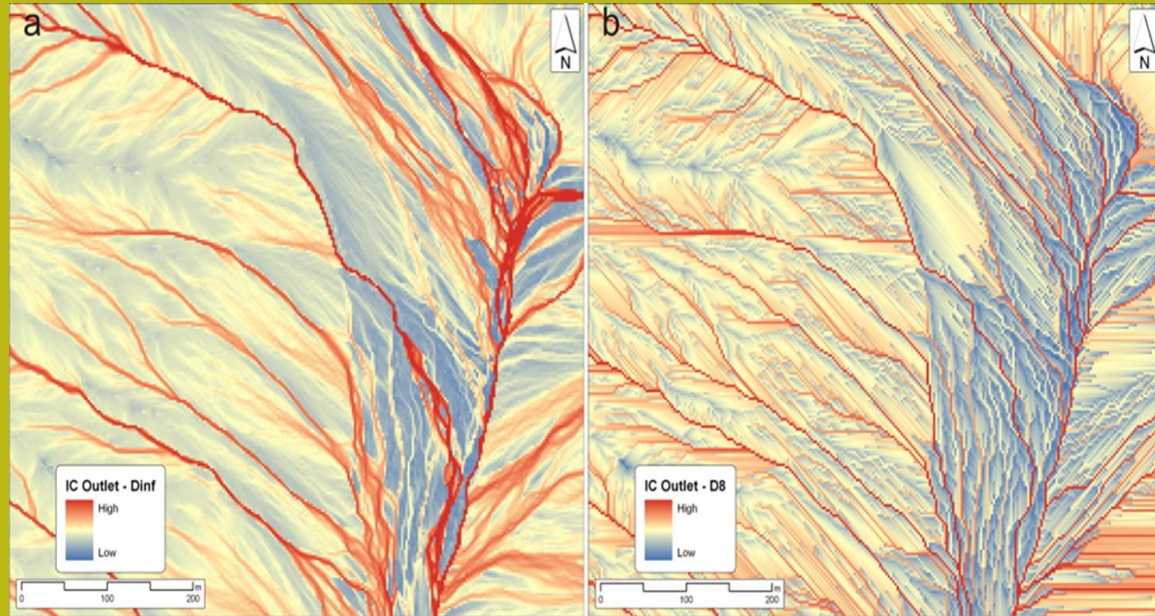
## Weighting factor $W$

- Related to the impedance to water and sediment fluxes;
- $C$  factor of USLE – RUSLE in the original model;
- Replaced by a roughness index (Cavalli et al., 2008).

High  $W$  : Low roughness and low impedance to fluxes

Low  $W$  : High roughness and high impedance to fluxes

**Application to debris flow and Surface landslide  
process .. But not only**

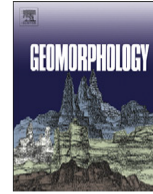


$$W = 1 - \left( \frac{RI}{MAX(RI)} \right)$$

$W$  (m/m) ranges from 0 to 1;  
minimum value set to 0.01

*standard dev of  
local residual*  
Cavalli (2008)

$$RI = \sqrt{\frac{\sum_{i=1}^{25} (x_i - x_m)^2}{25}}$$



**Site:** Büğdüz River catchment in SW Turkey (262 km<sup>2</sup>)

**DTM:** Detailed information not provided.

**aims:** to elucidate the spatial variability of sediment sources and geomorphic coupling between hillslope and floodplains

# A sediment fingerprinting approach to understand the geomorphic coupling in an eastern Mediterranean mountainous river catchment

Koen D'Haen, Bert Duser, Gert Verstraeten\*, Patrick Degryse, Hanne De Brue

Department of Earth and Environmental Sciences, KU Leuven, Celestijnenlaan 200E box 2409, B-3001 Leuven, Belgium  
Centre for Archaeological Science, KU Leuven, Celestijnenlaan 200E box 2408, B-3001 Leuven, Belgium

Cavalli's approach was also used by Authors

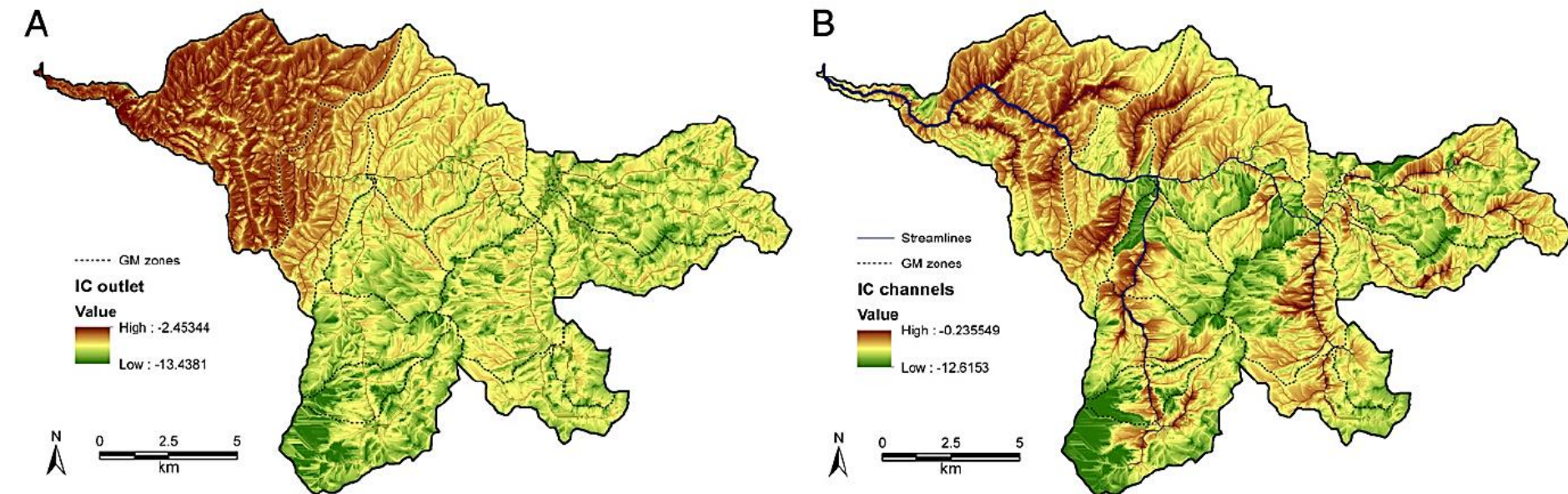


Fig. 4. (A) Connectivity index of Büğdüz catchment with respect to catchment outlet (IC outlet). (B) Connectivity index of Büğdüz catchment with respect to catchment outlet (IC channels).



# Initial hydro-geomorphic development and rill network evolution in an artificial catchment

Anna Schneider,<sup>1†\*</sup> Horst H. Gerke,<sup>2</sup> Thomas Maurer<sup>1</sup> and Rossen Nenov<sup>1</sup>

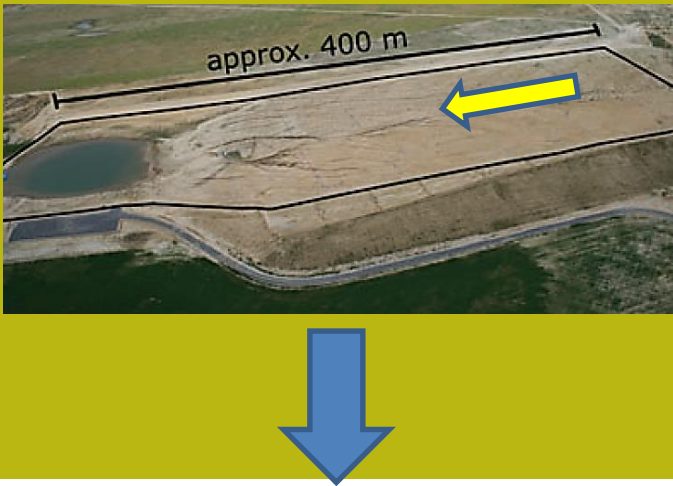
<sup>1</sup> Brandenburg University of Technology (BTU) – Research Centre for Landscape Development and Mining Landscapes (FZLB), Cottbus, Germany

<sup>2</sup> Leibniz-Centre for Agricultural Landscape Research (ZALF) – Institute of Soil Landscape Research, Müncheberg, Germany

Received 2 February 2012; Revised 18 December 2012; Accepted 20 December 2012

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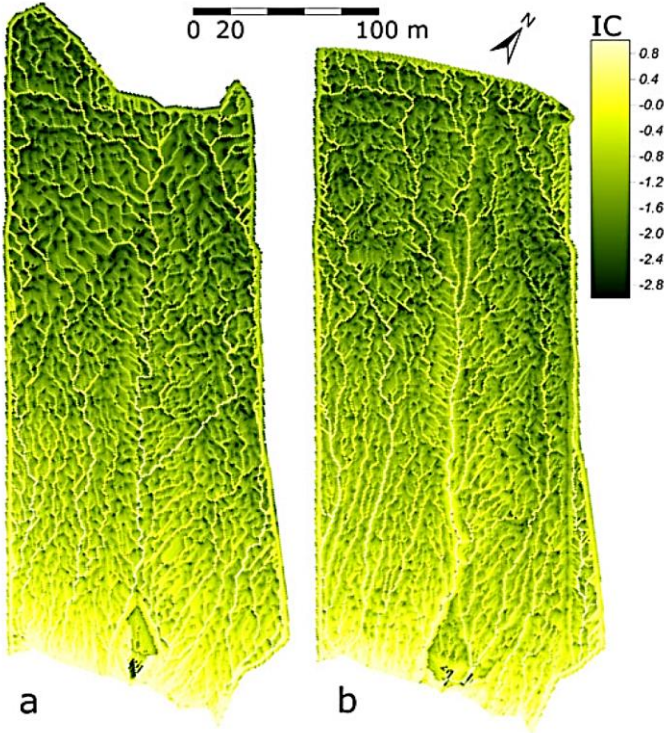
<sup>†</sup>Current affiliation: 2 Brandenburg University of Technology (BTU) - Chair of Geopedology and Landscape Development, Cottbus, Germany



**Site:** open-cast lignite mine Cottbus,  
150 km south of Berlin (Germany)  
Approx 0.04 km<sup>2</sup>

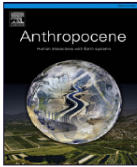
**DTM:** 1x1 m survey in several phases  
of erosion evolution (during 5 years)

**aims:** to characterize and to identify  
characteristic phases of rill network  
development in the artificially-  
created catchment as an example for  
initial hydrogeomorphic  
landform development in temperate  
climate



**Figure 13.** Spatial distribution of DEM cell values of the IC in the erosion-dominated area (as defined in Figure 3b), based on DEMs for (a) November 2005, and (b) March 2010. Grid cell size is 1 m by 1 m. Modification of the surface catchment area in autumn 2009 results in lower IC values in b).





**Site: Fukushima Prefecture**  
(Japan) Approx 600 km<sup>2</sup>

**DTM: 10x10 m**

**aims:** In post-accidental context, the paper aims to provide alternative **methods** to estimate the early dispersion of contaminated sediment during the 20 months that followed the nuclear accident in the mountainous catchments exposed to a succession of erosive rainfall, snowfall and snowmelt events.

## Tracking the early dispersion of contaminated sediment along rivers draining the Fukushima radioactive pollution plume

Caroline Chartin<sup>a</sup>, Olivier Evrard<sup>a,\*</sup>, Yuichi Onda<sup>\*\*,b</sup>, Jeremy Patin<sup>b</sup>, Irène Lefèvre<sup>a</sup>, Catherine Ottlé<sup>a</sup>, Sophie Ayrault<sup>a</sup>, Hugo Lepage<sup>a</sup>, Philippe Bonté<sup>a</sup>

<sup>a</sup> Laboratoire des Sciences du Climat et de l'Environnement (LSCE/IPSL), Unité Mixte de Recherche 8212 (CEA, CNRS, UVSQ), 91198 Gif-sur-Yvette Cedex, France

<sup>b</sup> Center for Research in Isotopes and Environmental Dynamics (CRIED), University of Tsukuba, 1-1-1 Tennodai, Tsukuba, Ibaraki 305-8572, Japan

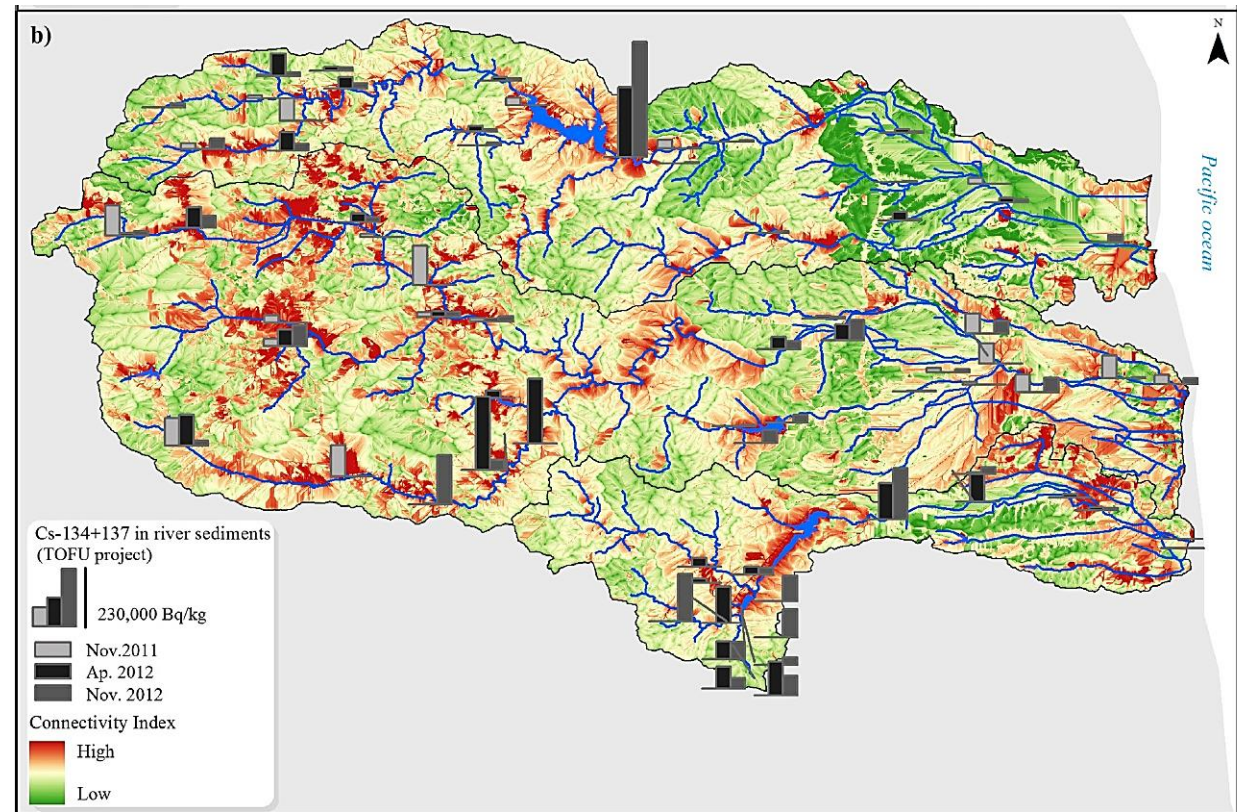


Fig. 6. Dominant land uses in the coastal catchments derived from analysis of satellite images (a) and associated hillslope-to-sinks hydro-sedimentary connectivity index compared to river sediment radiocaesium activities (b). ((M) Mano catchment); (N) Nitta catchment; (O) Ota catchment).

**Site:** in Swiss National Park (SNP), Engadine region in Switzerland - study area 6.4km<sup>2</sup>

**DTM:** LiDAR-based 2x2m

**aims:** 1) to evaluate the morphometric GIS modelling results against the field based geomorphic map, 2) to decipher key controls on the present-day sediment flux in a small, de-glaciated mountain valley, and 3) to address the question of whether traditional geomorphic field maps have become indispensable today when studying mountain cascading systems.

*Cavalli et al (2013) variant for W factor*

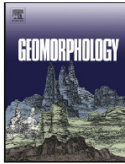


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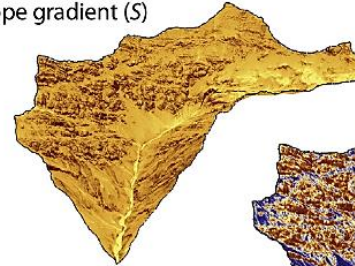


Sediment connectivity in the high-alpine valley of Val Mütsch, Swiss National Park – linking geomorphic field mapping with geomorphometric modelling

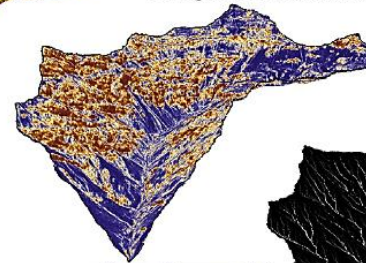
Karoline Messenzehl \*, Thomas Hoffmann, Richard Dikau

Department of Geography, University of Bonn, Meckenheimer Allee 166, 53115 Bonn, Germany

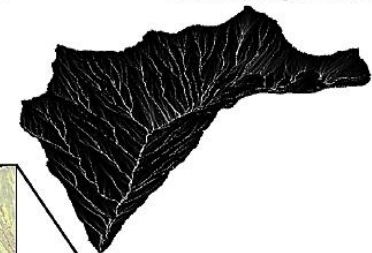
Slope gradient (S)



Roughness index (RI)



Contributing area (A)



Upslope component

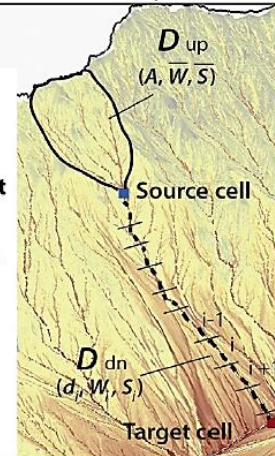
$$D_{up} = \overline{W} \overline{S} \sqrt{A}$$

Downslope component

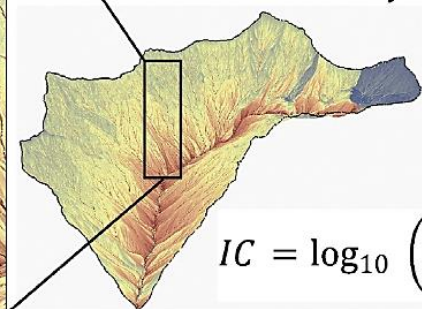
$$D_{dn} = \sum_i \frac{d_i}{W_i S_i}$$

Weighting factor W

$$W = 1 - \left( \frac{RI}{RI_{max}} \right)$$

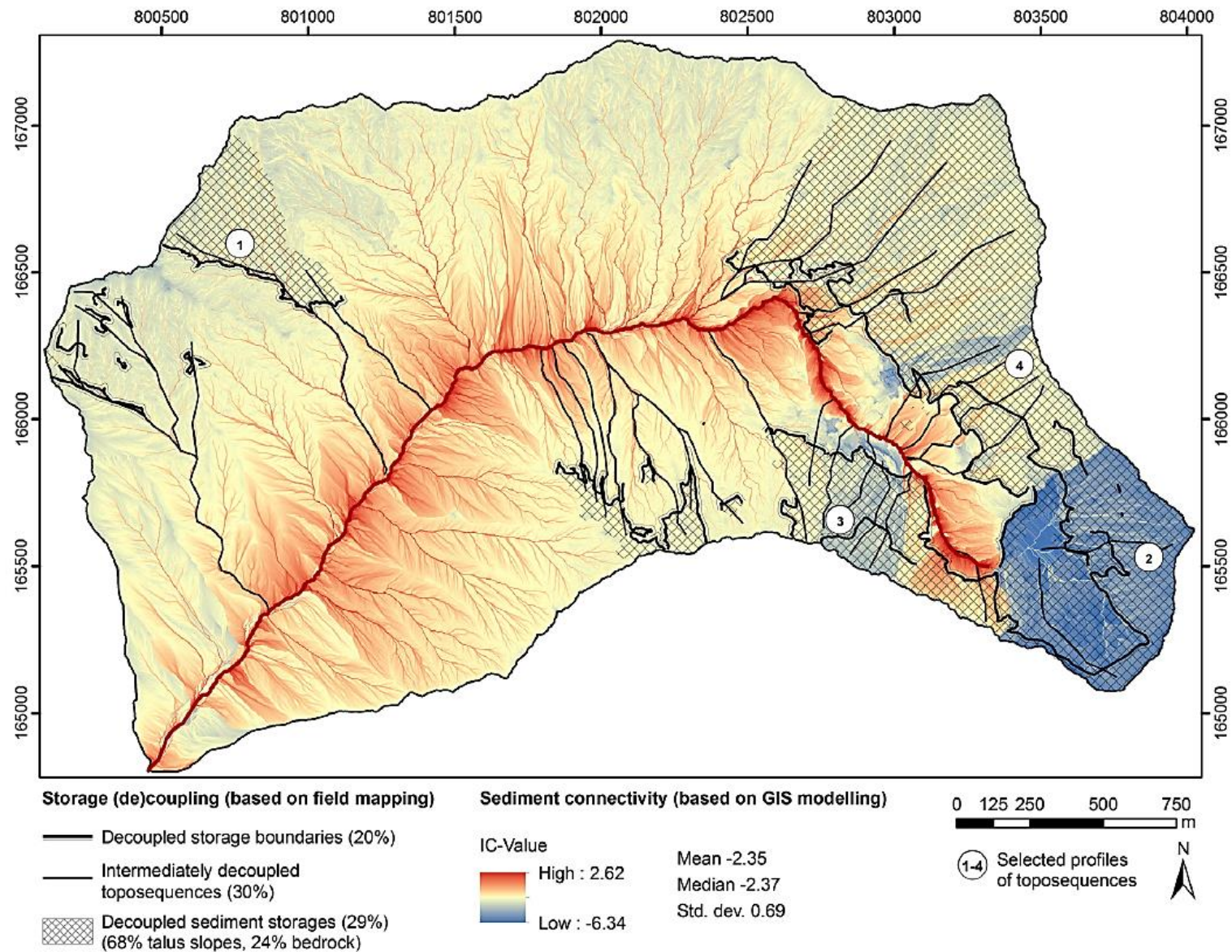


Index of connectivity (IC)



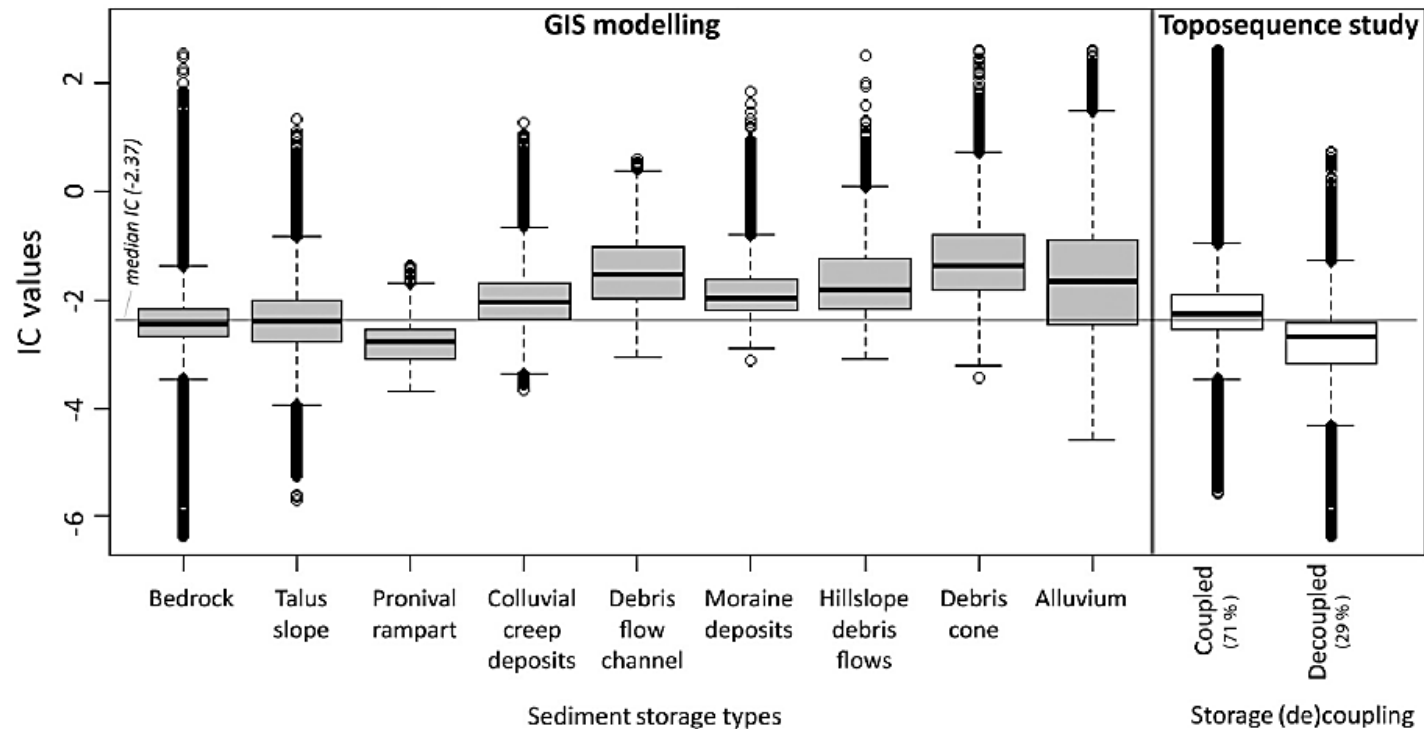
$$IC = \log_{10} \left( \frac{D_{up}}{D_{dn}} \right)$$





**Fig. 9.** Storage (de)coupling, according to the field-based analysis of toposequences, and index of connectivity (IC), calculated by the GIS modelling approach (2-m DTM, © SwissTopo, in 2013). Thick double lines show storage boundaries qualitatively defined as decoupled due to lacking sediment transfer between adjacent landforms caused by inactivity of geomorphic processes or the occurrence of buffers. As a consequence, around 29% of the basin surface has no connectivity to the fluvial system (crosshatched area). Among them, 68% of the talus slopes and 24% of the bedrock-coverage are affected by this disconnection.

## IC related to different geomorphological units and processes depending on sediment storage types



**Fig. 10.** IC values of sediment storage types in Val Mütsch and comparison between the modelling results (grey boxplots, left-hand side) and the field-based mapping results of the toposquence study (white boxplots right-hand side). Around 35% of the basin surface, which has been qualitatively classified as being decoupled, is related to IC values higher than the basin's median IC.



# Distributed empirical algorithms to estimate catchment scale sediment connectivity and yield in a subtropical region

Reza Jamshidi,<sup>1\*</sup> Deirdre Dragovich<sup>1</sup> and Ashley A. Webb<sup>2,3</sup>

<sup>1</sup> School of Geosciences F09, The University of Sydney, NSW, 2006, Australia

<sup>2</sup> Forestry Corporation of NSW, PO Box 4019, Coffs Harbour Jetty, NSW, 2450, Australia

<sup>3</sup> Australian Centre for Agriculture and Law, University of New England, NSW, 2351, Australia

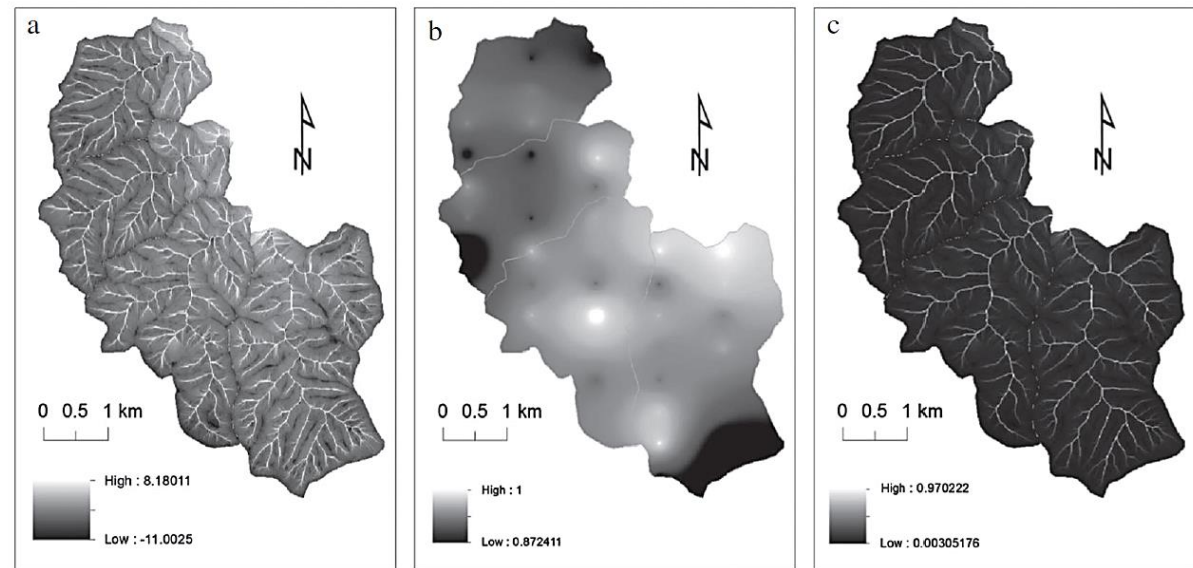


Figure 2. The predicted maps of (a) index of connectivity IC, (b) the maximum theoretical SDR coefficient  $SDR_{max}$  and (c) SDR variability within the study area. The IC and SDR maps in this figure were selected for spatial variability in 2007. Catchment boundaries are depicted in Figure 1

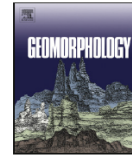
$$SDR = \frac{SDR_{Max}}{1 + \exp\left(\frac{Ic_0 - Ic}{k}\right)}$$

Use of SDR by IC, and optimized  $IC_0$  and  $k$ , by Vigiak et al. 2012  
 $IC_0=0.5$  and  $k=2.0$

**Site:** Kangaroo River  
 State forest, northern  
 NSW, Australia  
 21.7 km<sup>2</sup>.

**DTM:**10x10 m

**aims:** to apply a  
**distributed hillslope  
 erosion-SDR  
 approach** in raster  
 data layers **to assess  
 the impacts of  
 vegetation removal  
 (single tree selection  
 logging) on the  
 spatial distribution  
 of  
 estimated sediment  
 yields**



## Connectivity structure of the Kosi megafan and role of rail-road transport network

Rakesh Kumar <sup>a</sup>, Vikrant Jain <sup>a,b,\*</sup>, G. Prasad Babu <sup>c</sup>, Rajiv Sinha <sup>d</sup>

<sup>a</sup> Department of Geology, Centre of Advanced Studies, University of Delhi, Delhi 110007, India

<sup>b</sup> Discipline of Earth Sciences, Indian Institute of Technology Gandhinagar, Ahmedabad, Gujarat 382424, India

<sup>c</sup> UNDP, Disaster Management Unit, 55 Lodi Estate, New Delhi 110003, India

<sup>d</sup> Department of Earth Sciences, Indian Institute of Technology Kanpur, Kanpur 208 011

**Site:** Kosi megafan, India  
Hymalayan region.  
approx. 7000 km<sup>2</sup>

**DTM:** SRTM DEM data of  
February 2002.

<http://www.cgiar-csi.org>  
(resolution not provided)

**Aims:** The paper presents  
the **two-dimensional**  
**dis(connectivity)** structure  
of the Kosi megafan, India,  
including the lateral and  
longitudinal dimensions  
(continued ...)

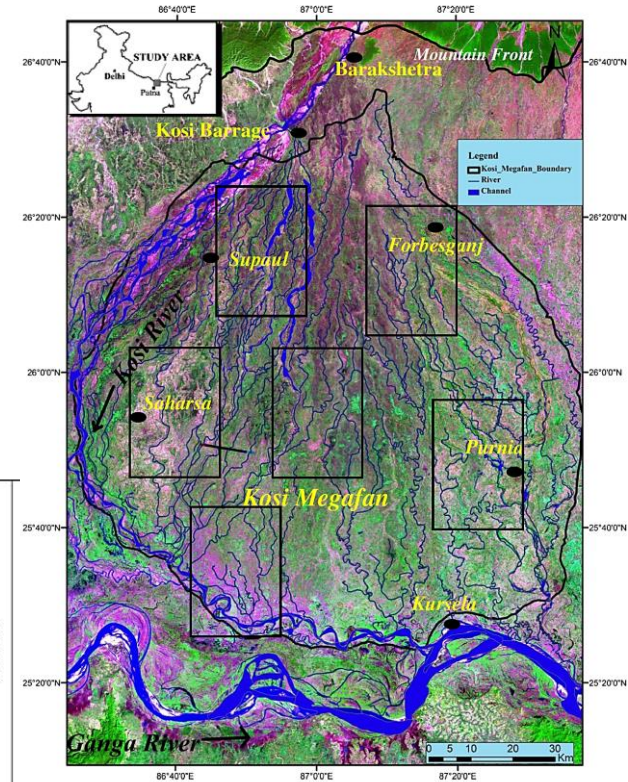
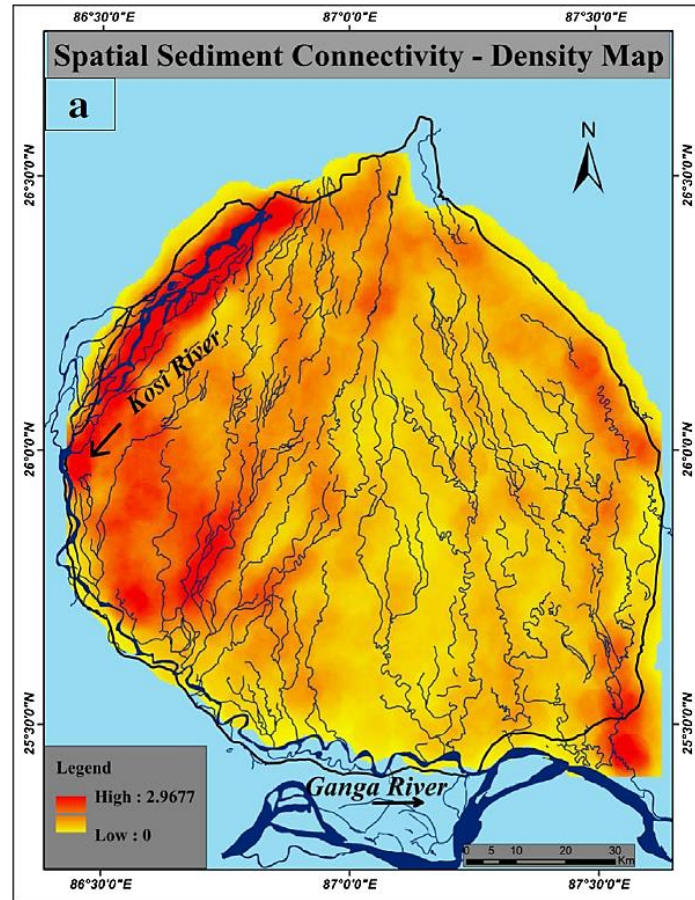


Fig. 1. Study area showing the Kosi River in the North Bihar, India. The megafan surface has numerous small paleochannels most of which are activated during the monsoon period. The Kosi River debouches into the plains at Barakshetra and joins the Ganga River system at Karsela. Windows for field-based classification of the intersection points have been marked.

(continued ...) of  
geomorphic connectivity.  
The quantitatively  
defined and the  
'anthropogenic' impacts  
on the connectivity  
structure due to railroad  
transport network were  
also estimated.

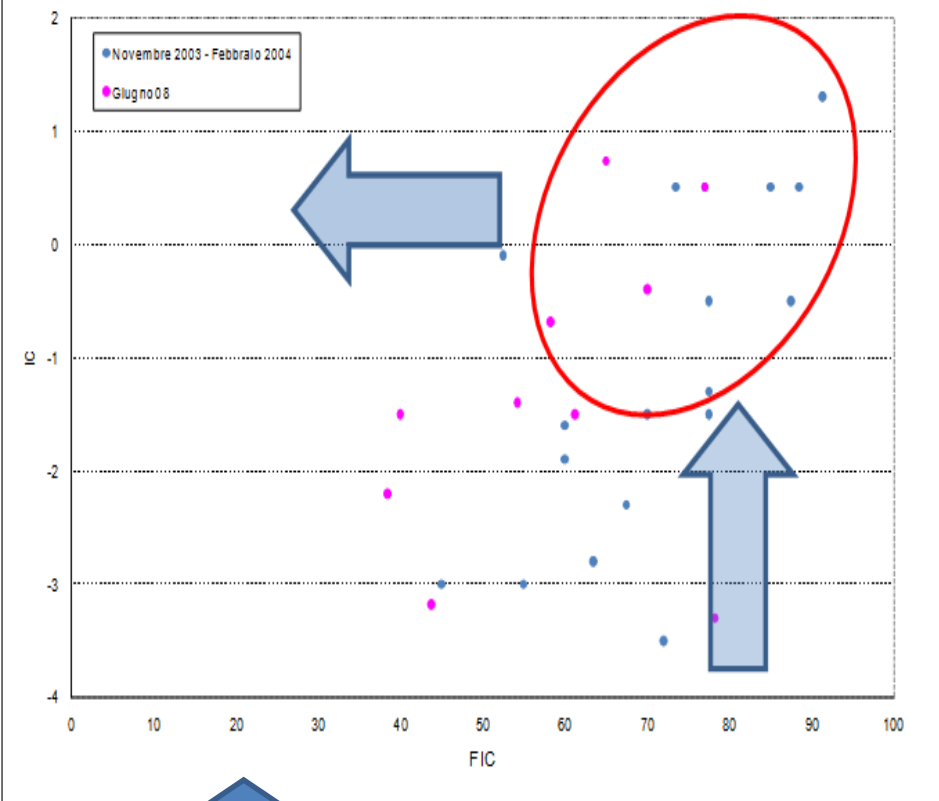
# IC computations and variants

## Some main tested variants:

- Areas from 0.04 km<sup>2</sup> to 7000km<sup>2</sup>
- DEM resolution: from 1x1 m to 50X50m
- Use (or not) of river mask as local sink
- Variable Minimum contributing area to generate permanent drainage lines (river mask) from 0.5 ha to 20 ha, or more...
- Use (or not) of roads and urban areas as SINK (like river mask do)
- W factor ( classical is  $W=C(\text{USLE TYPE})$  ) or  $W=RI$  (Cavalli 2013) or  $W=ER$  (Sougnez et al. 2011)
- Different countries, environments, landscapes , climates and anthropogenic impacts
- At moment has been published 11 applications



Confronto indice FIC e IC

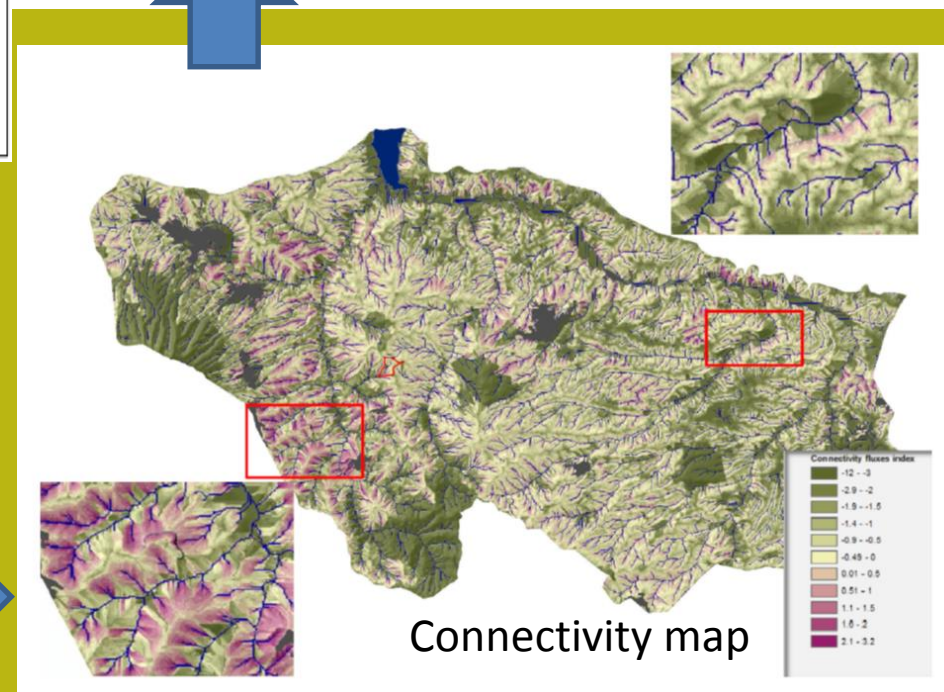
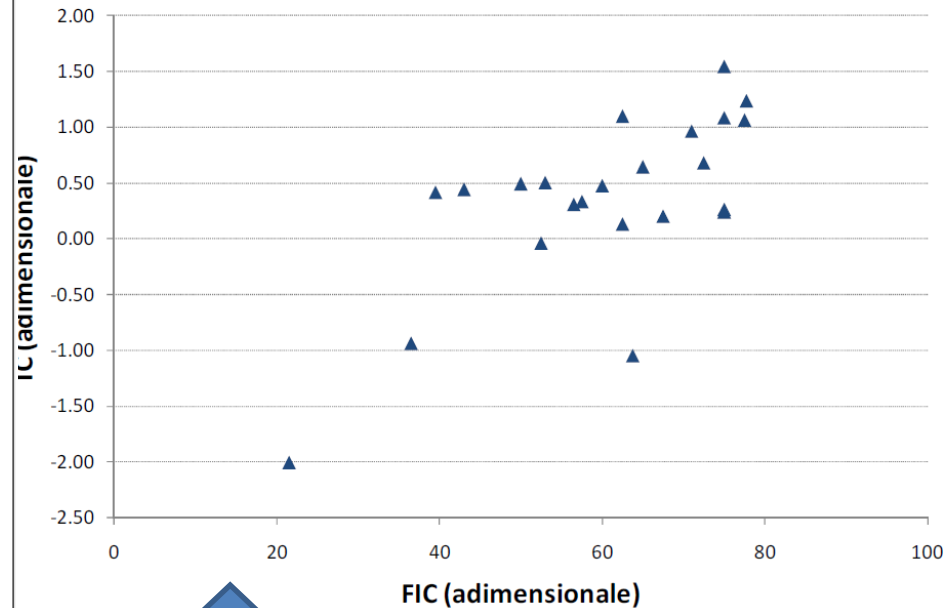


Extended dataset IC versus FIC  
(Bilancino watershed; Cassi, 2010)

In both cases  
we observe high  
connectivity for  
 $IC > 0.0$

Rendina watershed,  
South Italy (240 km<sup>2</sup>)  
DTM 20x20  
Project DESIRE

Confronto fra FIC e IC



Connectivity map



# W factor: an evolution for IC2 model

The second version of IC accounts of : *hydrologic soil properties , magnitude of rainfall event, surface roughness*

$$C_r = 1 - \frac{I_{dt} + S}{P_{dt}}$$

Step 1: computation event runoff coefficient ( $C_r$ ) (adimensional) , by total infiltration volume ( $I_{dt}$ ), surface water storage ( $S$ ) and rainfall volumes ( $P_{dt}$ ) (all in mm)  
Infiltration can be calculated locally by model e.g. model di Morel-Seytuox (1978) for each land units,  
 $S$  was computed with relationship by Borselli and Torri (2010, Journal of Hydrology), as a function of surface roughness and local slope gradient

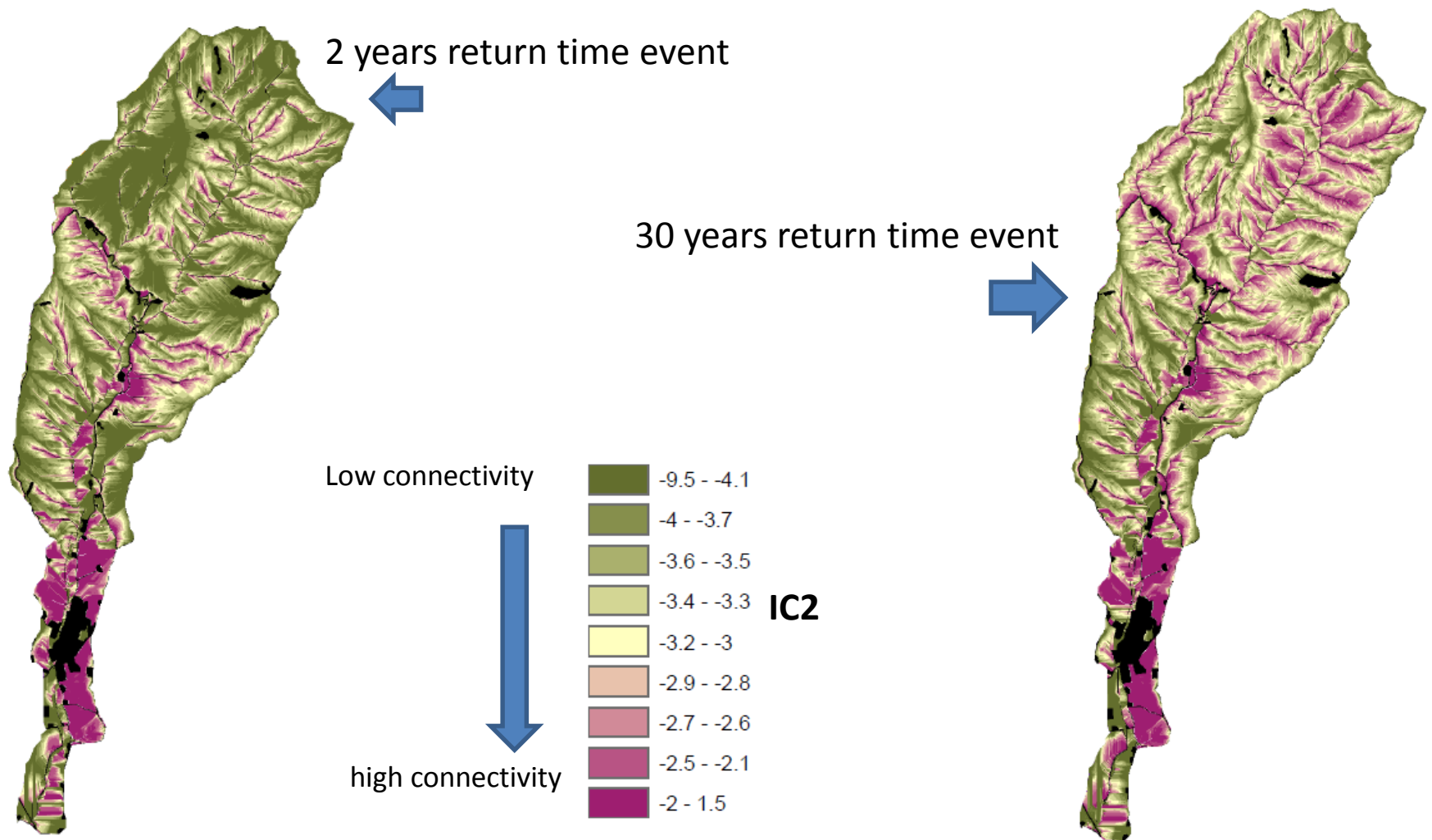
$$\frac{1}{f}$$

Step 2 Hydraulic roughness due land use and soil surface characteristics (including vegetation) by Darcy Weisbach ( $f$ ) friction factor (adimensional)

$$W = C_r * \frac{1}{f}$$

Step 3: Final  $W$  calculation  
(please note  $W$  is still adimensional )

Cassi, 2010 (PhD thesis)



Average Intensity 15.7 mm/h  
Duration 2.5 h  
Amount 39.2 mm

**First application of IC2  
Portion of Bilancino watershed**

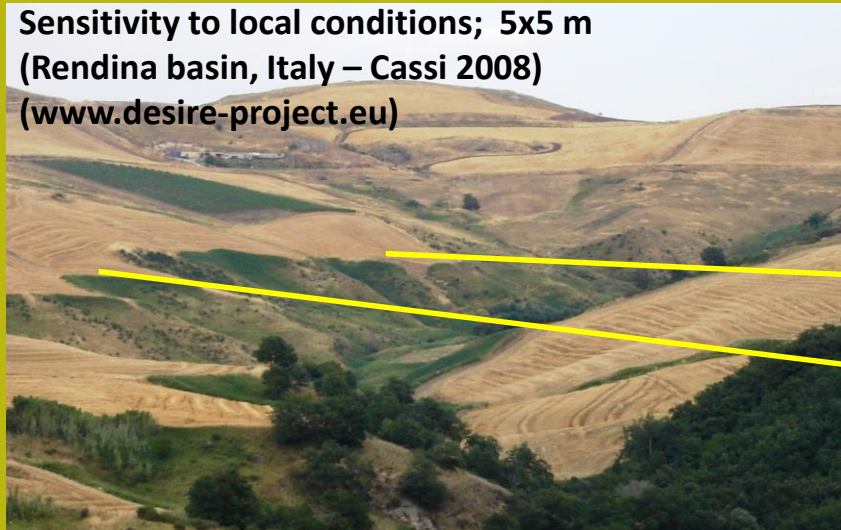
Average Intensity 37.9 mm/h  
Duration 1 h  
Amount 37.9 mm

Cassi, 2010 (PhD thesis)

## Scenario analysis of connectivity evolution due to land management practice

Prevailing land use of the watershed is wheat crop. The connectivity index is evaluated before and after harvesting (June) when the borders of the fields are ploughed for 5 meters wide.

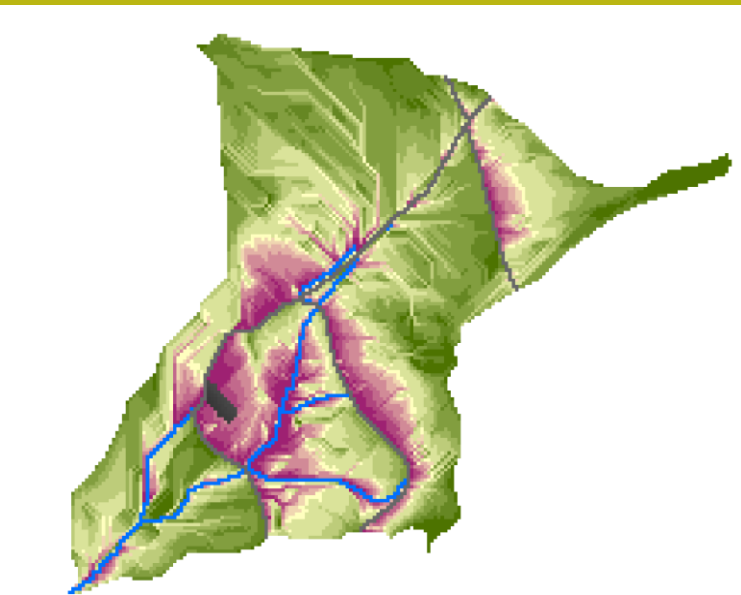
Sensitivity to local conditions; 5x5 m  
(Rendina basin, Italy – Cassi 2008)  
([www.desire-project.eu](http://www.desire-project.eu))



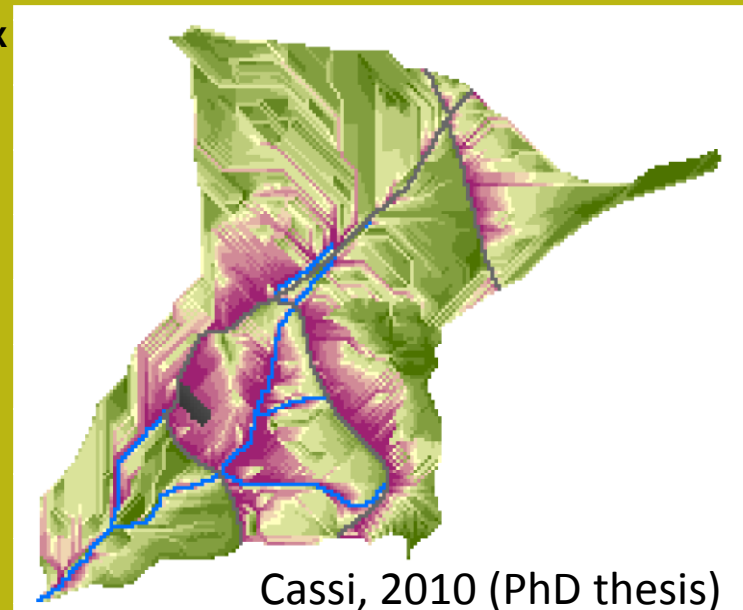
Ploughed field border with erosion evidences



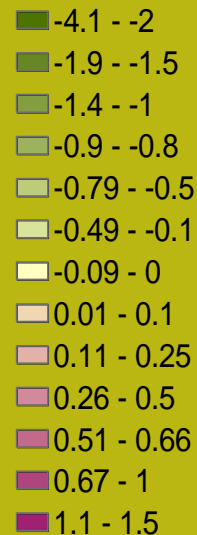
Without strip of bared soil on field border



With strip of bared soil on field border



Connectivity fluxes index



Cassi, 2010 (PhD thesis)





## Now .....Connectivity in Mexico...

Project MOPRI (2013-2015)

*“Modelado de procesos hidrológico, dinámica de hidrofobicidad e infiltración, para su aplicación en la evaluaciones del riesgo debido a inundaciones y lahars: aplicación en la ciudad de San Luis Potosì y en el Volcán De Colima” (2013-2015)(CONACYT-Ciencia Basica-2012-01 -184060)*

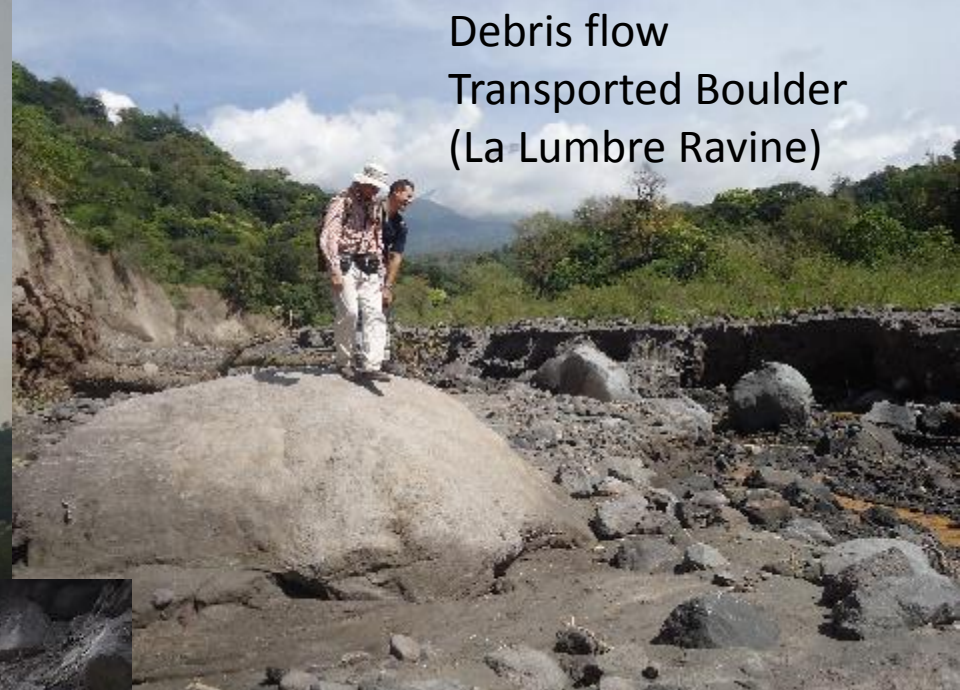
**Main geomorphic processes: pyroclastic flow, debris flows (lahars), mass movement, gully erosion, edifice collapse and debris avalanches, Approx. every 2500 years in the last 15000BP**







Colima,  
Volcan de  
Fuego



Debris flow  
Transported Boulder  
(La Lumbre Ravine)



Debris flow  
Transported Boulders  
(Montegrando ravine)



Infiltration Test  
(Montegrando Ravine)





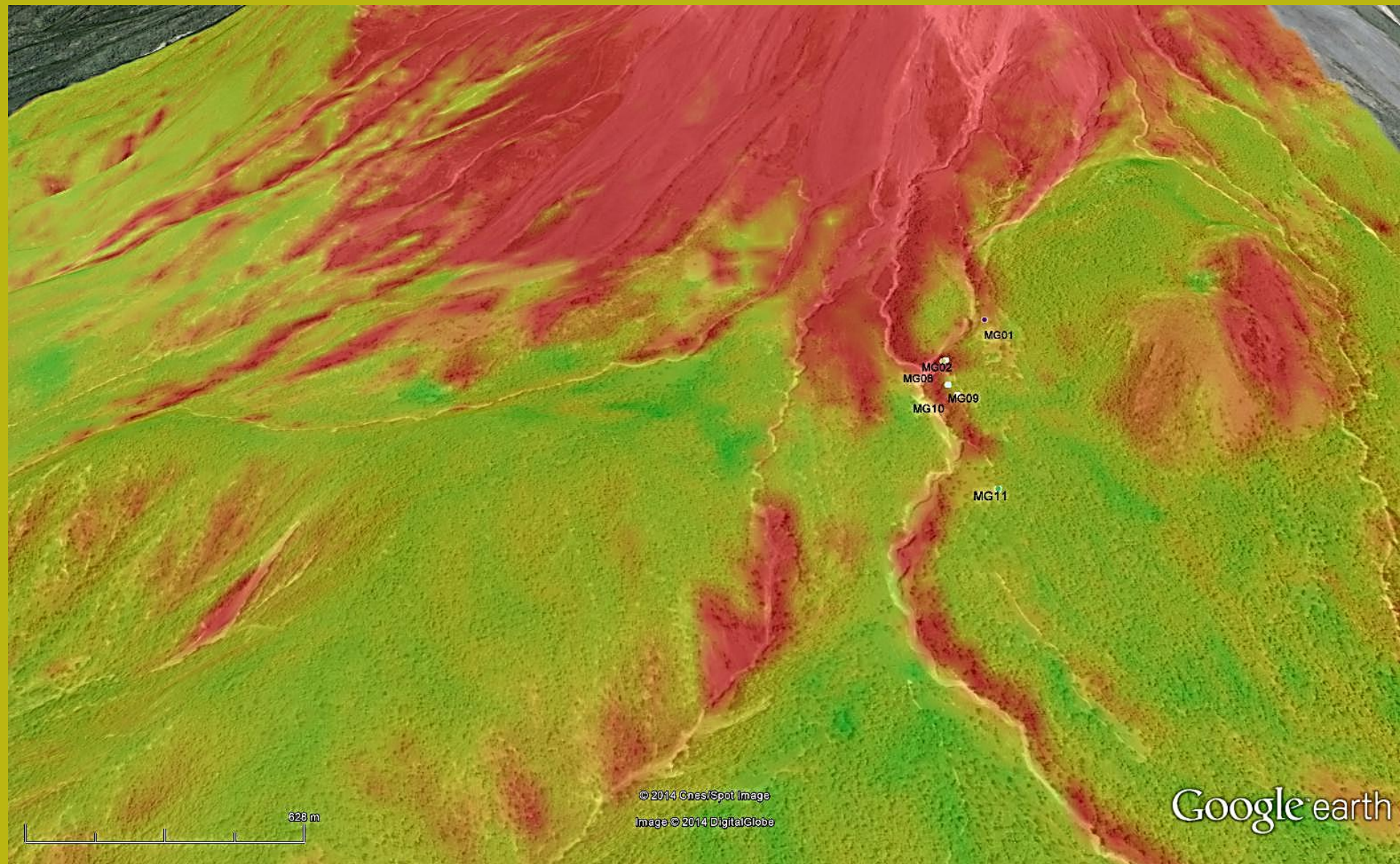
Connectivity map (IC) DEM 5x5m (transparent overlay on Google Earth image 2014)





Portion of Colima volcano South flank with several barrancas and main source areas of lahars (above 2500 m a.s.l.); local remobilization, mass movements, and instability of older deposits are in evidence.





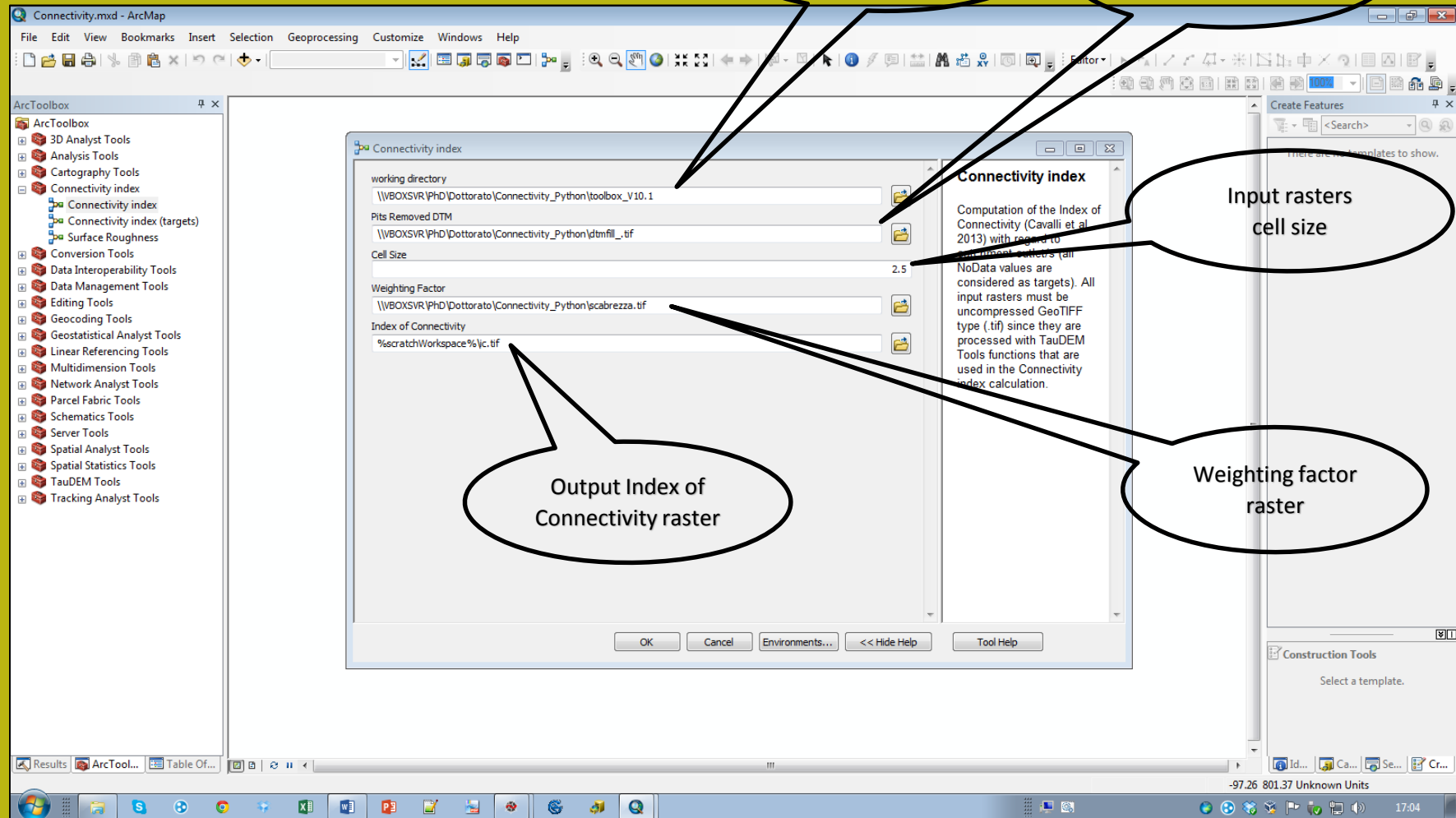
Connectivity may help understanding the dynamic of lateral hydrologic contributions that trigger or remobilize the lahars at the beginning of the rain season, when the soil is hydrophobic (Capra , Borselli et al. 2010).



# Connectivity: toolbox, software and extensions

# Cavalli's TOOLBOX (2014) for IC index computation

## ArcGIS implementation



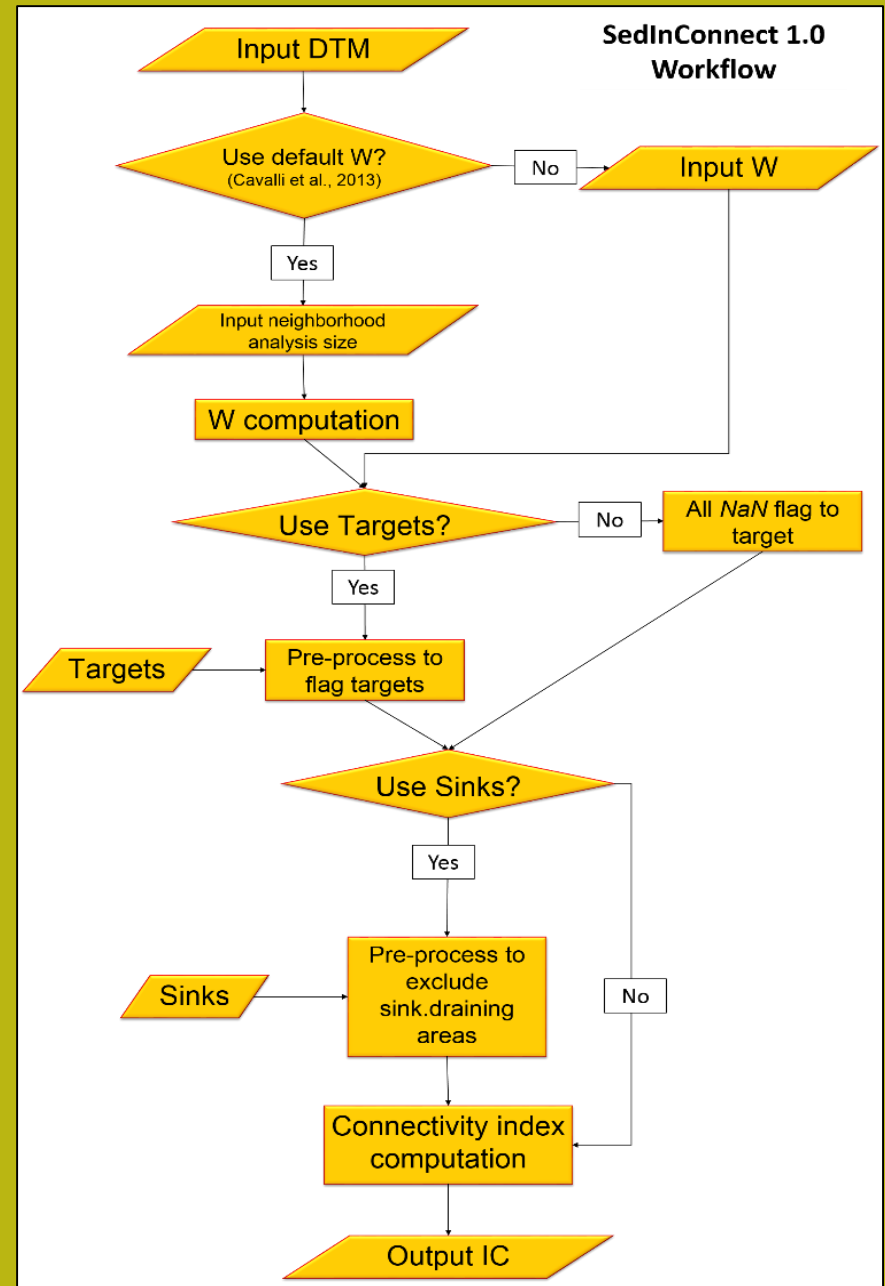
- Running under ArcGIS version 10.1 (with SP1!) and 10.2
- It requires the installation of TauDEM 5.1  
<http://hydrology.usu.edu/taudem/taudem5/downloads.html>

**Marco Cavalli 2014**  
marco.cavalli@irpi.cnr.it

# Stand-alone application (SedInConnect 1.0)

Marco Cavalli 2014  
marco.cavalli@irpi.cnr.it

- **Free application and open source code** - > thus encouraging its usage and receiving improvements from the scientific community.
- **Development based on the Python Experience** -> exploiting available libraries for efficient analysis of geographic datasets and guaranteeing cross-platform portability (Windows/Linux/MacOS), thus enhancing the potential users' community.
- **Stand-alone application** -> independent from any GIS software and more accessible to management authorities less involved in research issues but very close to management issues and priorities.

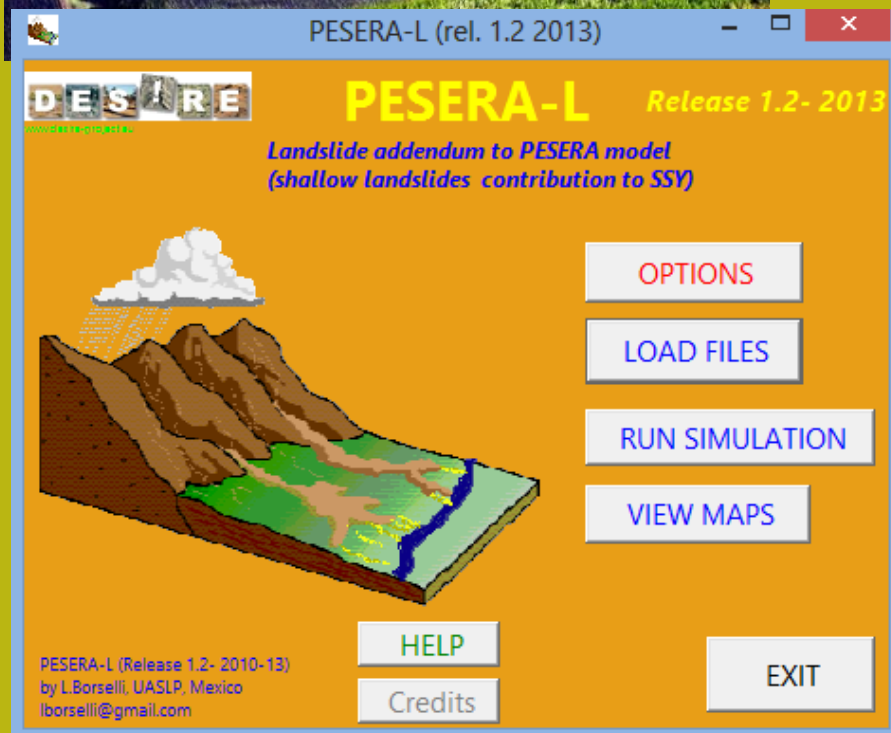






## ***PESERA-L, the shallow landslides contribution to specific sediment yield (SSY), as extensions of the PESERA soil erosion model***

See [www.Lorenzo-borselli.eu/pesera](http://www.Lorenzo-borselli.eu/pesera)  
For software download and documentation



# The Sediment delivery ratio for landslides SDR<sub>L</sub> And how to obtain SSY (under Pesera-L)

$$V = 10^6 A D \Psi SDR_L$$

$$SSY = \frac{V \gamma_s}{100A \Delta_t} \quad [Mg \ ha^{-1} yr^{-1}]$$

Where

***V***= net eroded Volume (m<sup>3</sup>)

***A***= area of HLU (km<sup>2</sup>)

***D***= average depth of landslides (m)

***Ψ***= fraction of area potentially unstable (-)

***SDR<sub>L</sub>***= sediment delivery ratio from landslides (-)

***γ<sub>s</sub>***= soil unit weight (Mg/m<sup>3</sup>)

***Δ<sub>t</sub>***= annual frequency (yr)

***SSY***= specific sediment yield from hillslope [Mg/ha/yr]





# Exponential distribution model for sediment delivery

Derived by  
*Miller and Burnett (2008)* and  
modified by implementation of  
a portion of IC

Where:

$L_R$  = landslide average runout (m)

$D_{dn}$  = Downslope routing  
weighted distance (m)

(downslope component IC model Borselli et al. 2008)

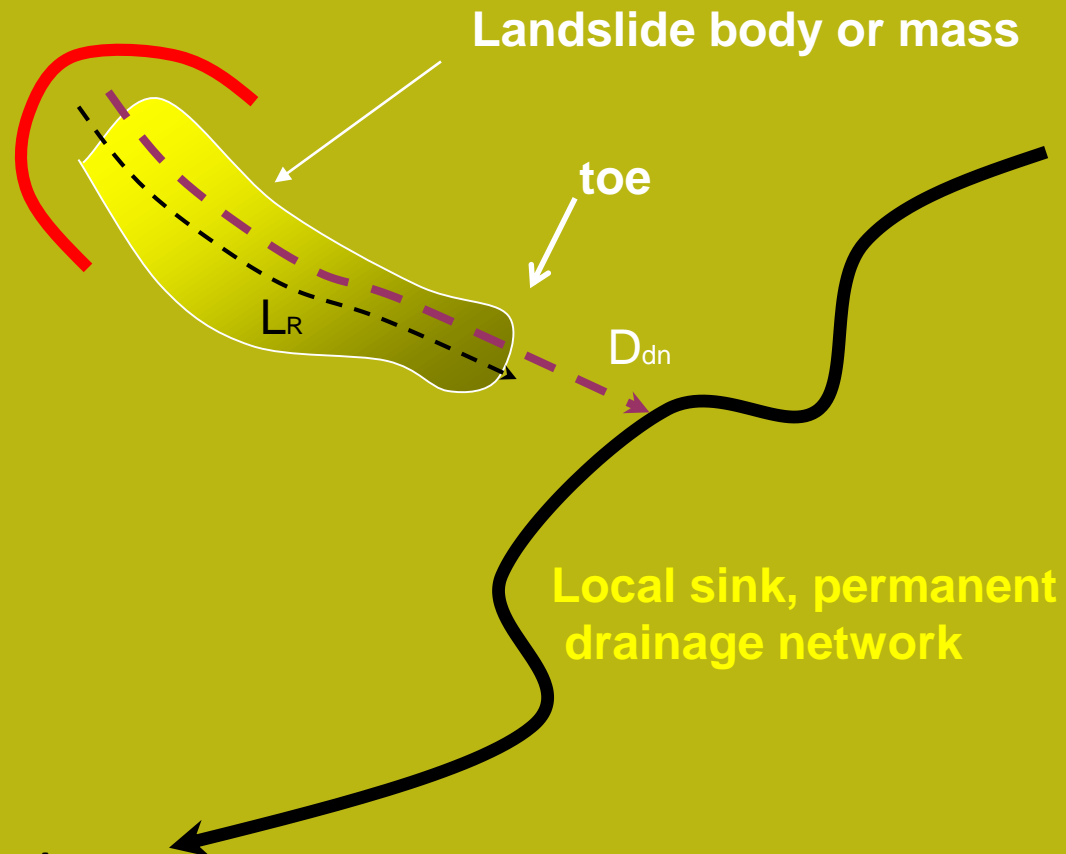
$$SDR_L = e^{-\lambda D_{dn}}$$

$$\lambda = \frac{1}{\bar{L}_R}$$

$$SDR_L = e^{-\frac{D_{dn}}{\bar{L}_R}}$$

*Mobility parameter*

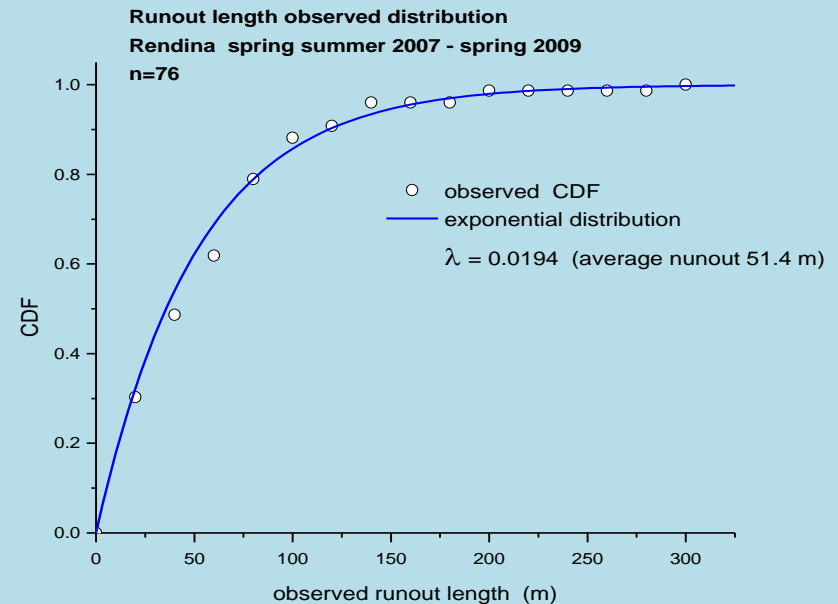
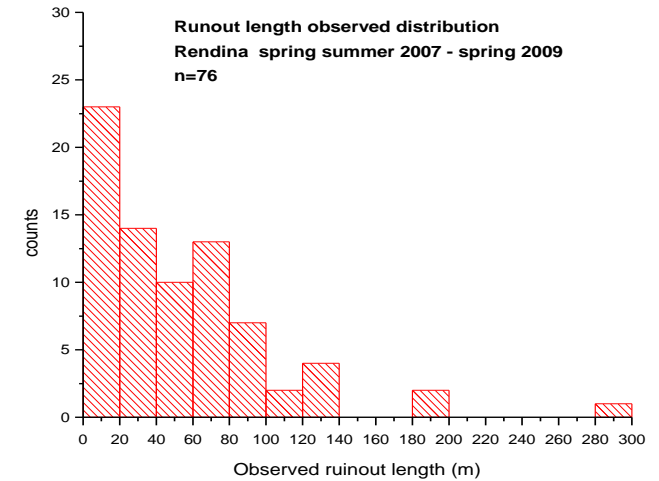
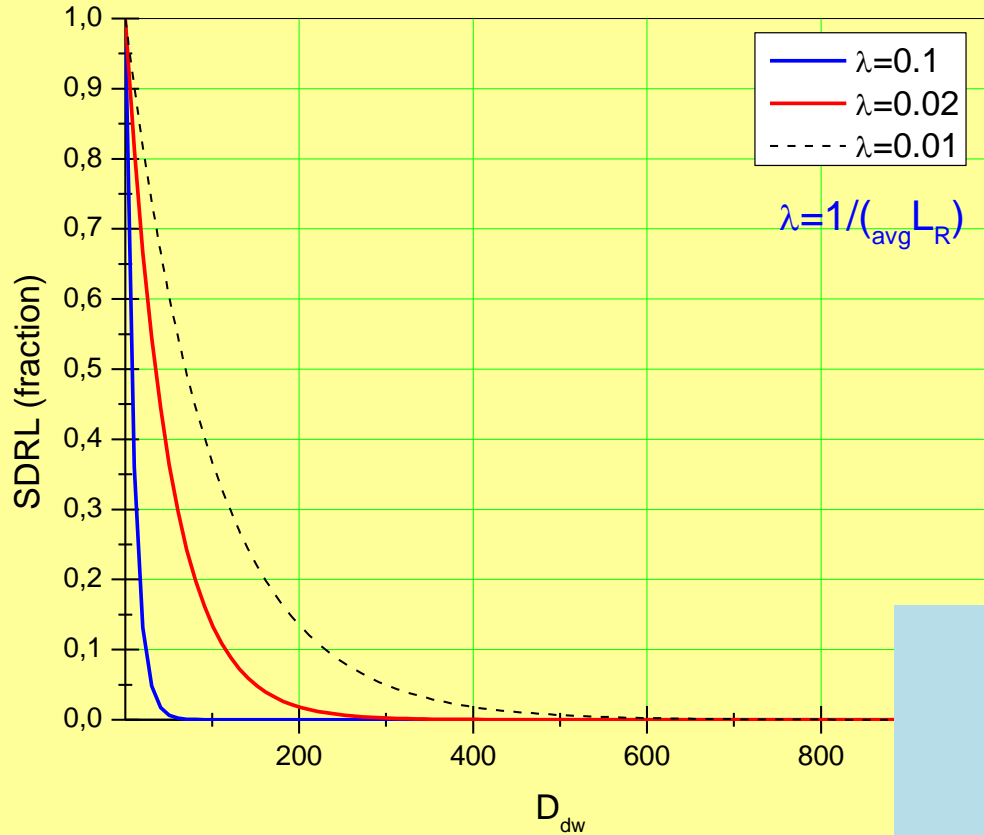
*that depends on average observed runout  $L_R$  and local  $D_{dn}$*



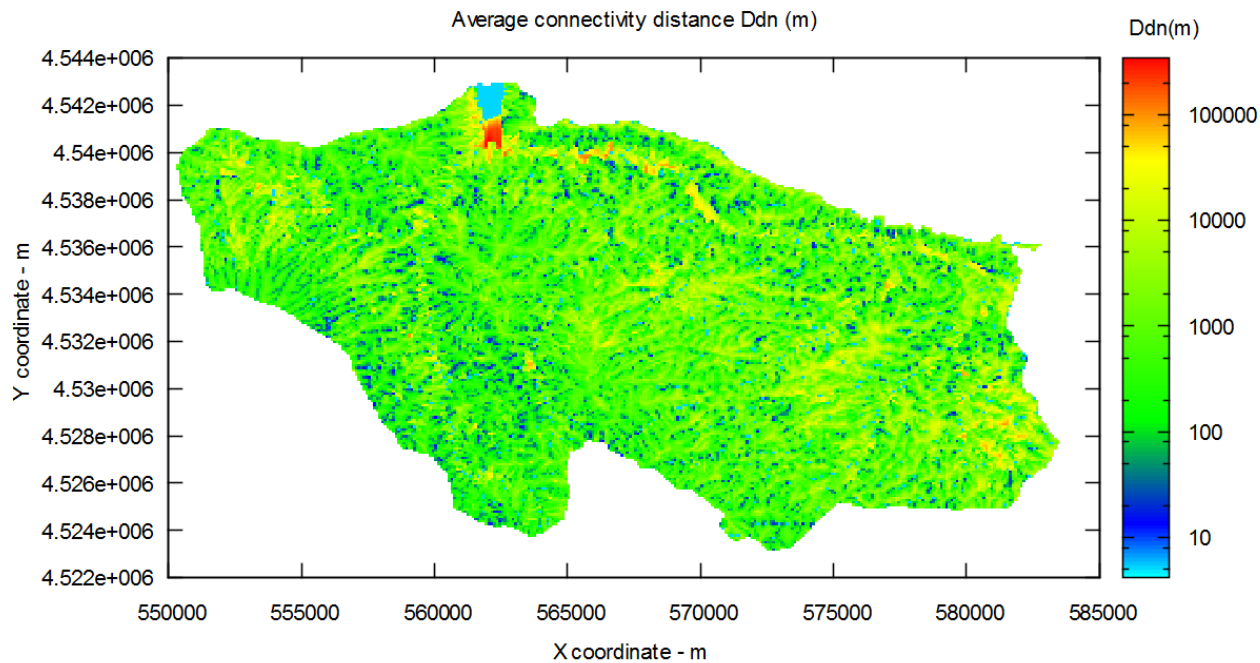


# probabilistic model of landslides and debris flow delivery to stream channels

( Miller & Burnett, 2008)



**Exponential probability distribution function depends from the average runout length  $L_r$  (measured) and the local site  $D_{dn}$  distance to a sink**

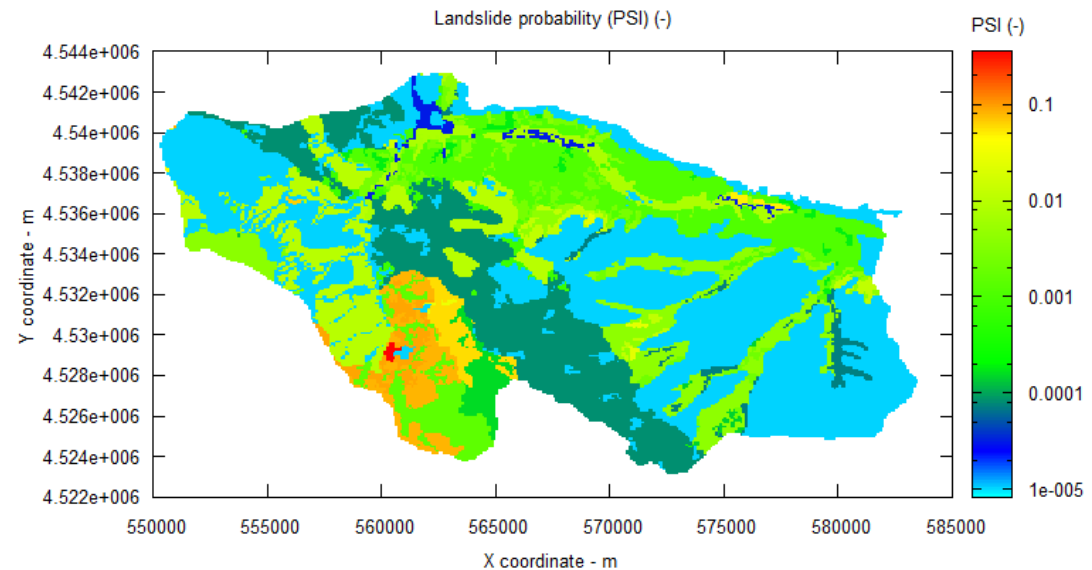


Application to Rendina  
watershed  
Project DESIRE

PESERA-L rel. 1.2 (2013) by L. Borselli, lborselli@gmail.com  
<http://www.lorenzo-borselli.eu/peseral>

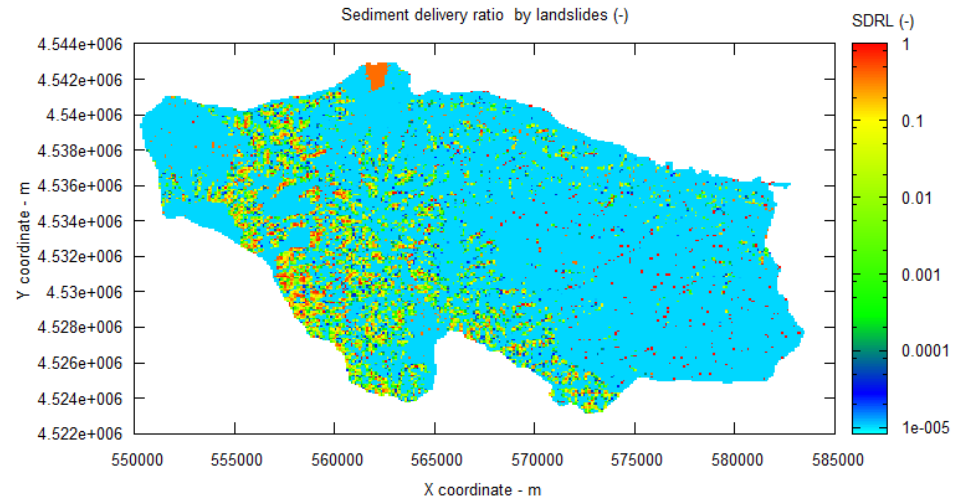
Connectivity average  
downslope distance

Landslide probability

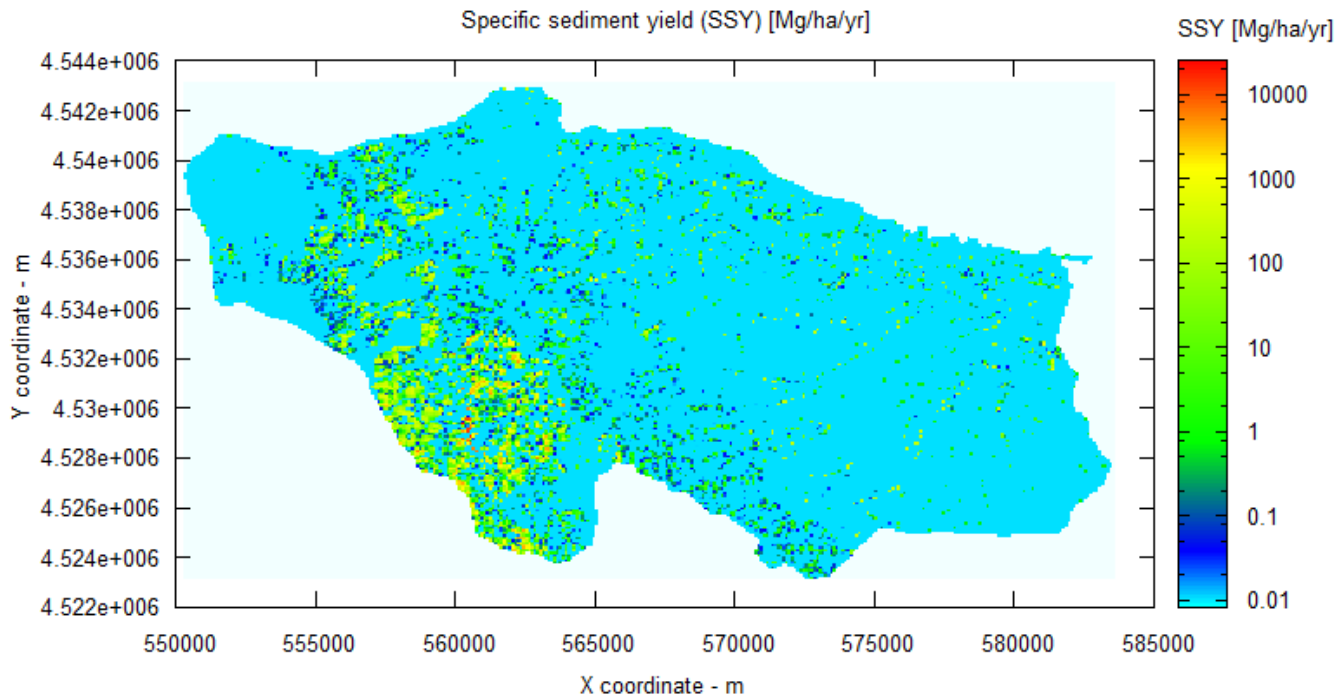


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# SDRL Sediment delivery by Shallow landslides



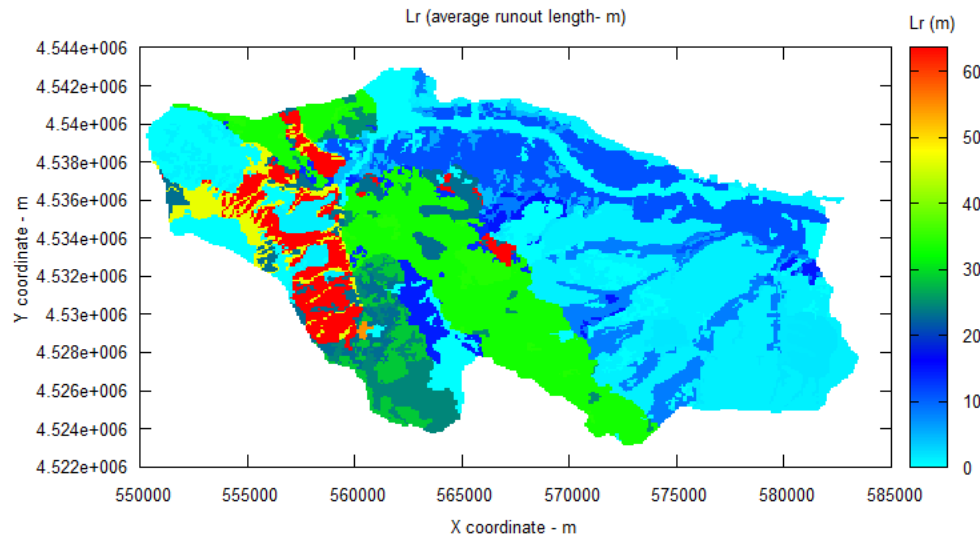
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SSY by landslides

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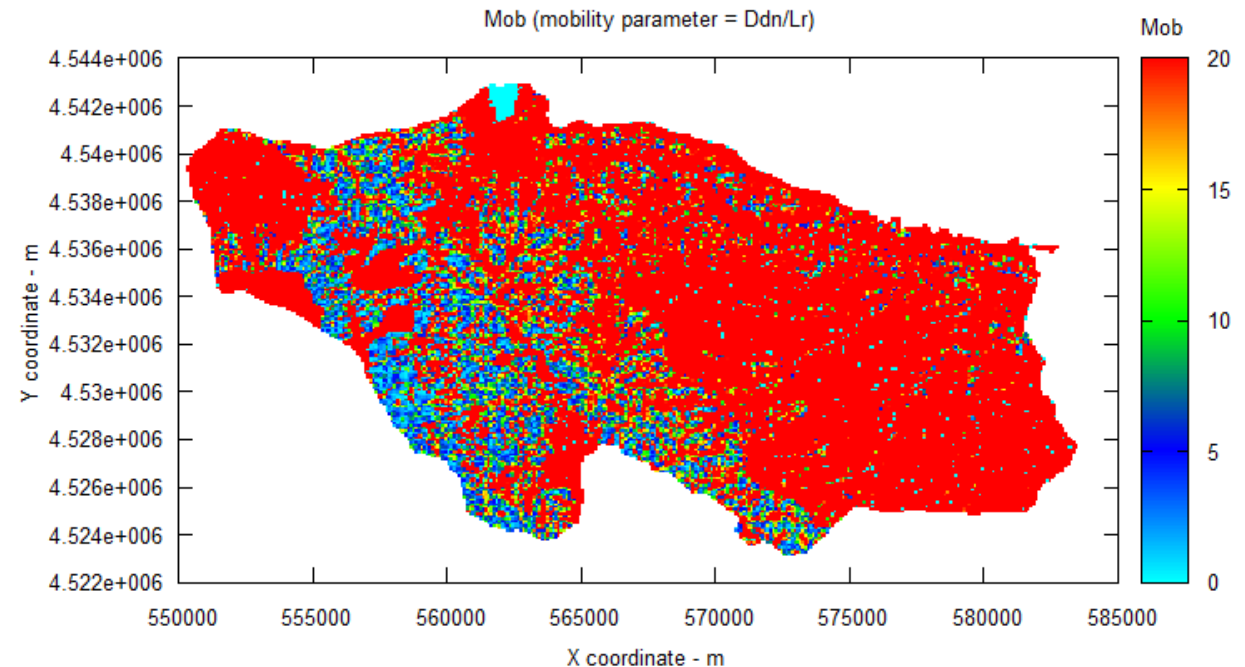


Average run-out length of shallow landslides for each LUS (land unit system) obtained from field survey and multi-temporal aerial photos

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Mobily parameter of landslides

$$\frac{\overline{D_{dn}}}{\overline{L_R}}$$



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**Mass  
movement  
type**



**Flow slide  
mudflow**

**Shallow  
Translational**

**Shallow  
Rotational**

**creeping**

$$\frac{\overline{D}_{dn}}{\overline{L}_R}$$

$$\overline{L}_R$$



**Land unit  
landforms**

**Badlands  
Clay shale  
Deposits  
High drainage  
density**

**rolling topography  
Medium steepness  
and medium drainage  
density**

**Rolling to flat  
topography**



# Landslide mobility parameter

*And the possible dependence from processes and landforms*

# Highlights and speculations - 1

The Connectivity index IC and FCA **provide an estimate of the potential connection index between the sediment eroded from hillslopes and the stream system or other local sinks;**

FCA can put in relation the IC index and SDR. SDR can be used then to correct the USLE-TYPE models (*transport capacity unlimited*) generally used for large catchments modelling and obtain a Sediment yield assessment .

The use of SDRmax limited values (Borselli et al. 2009 and Vigiak et al. 2012) can correct for inconsistencies indicated by some researchers (Kinnell 2004, Parson et al. 2006) in SDR assessment.



## Highlights and speculations - 2

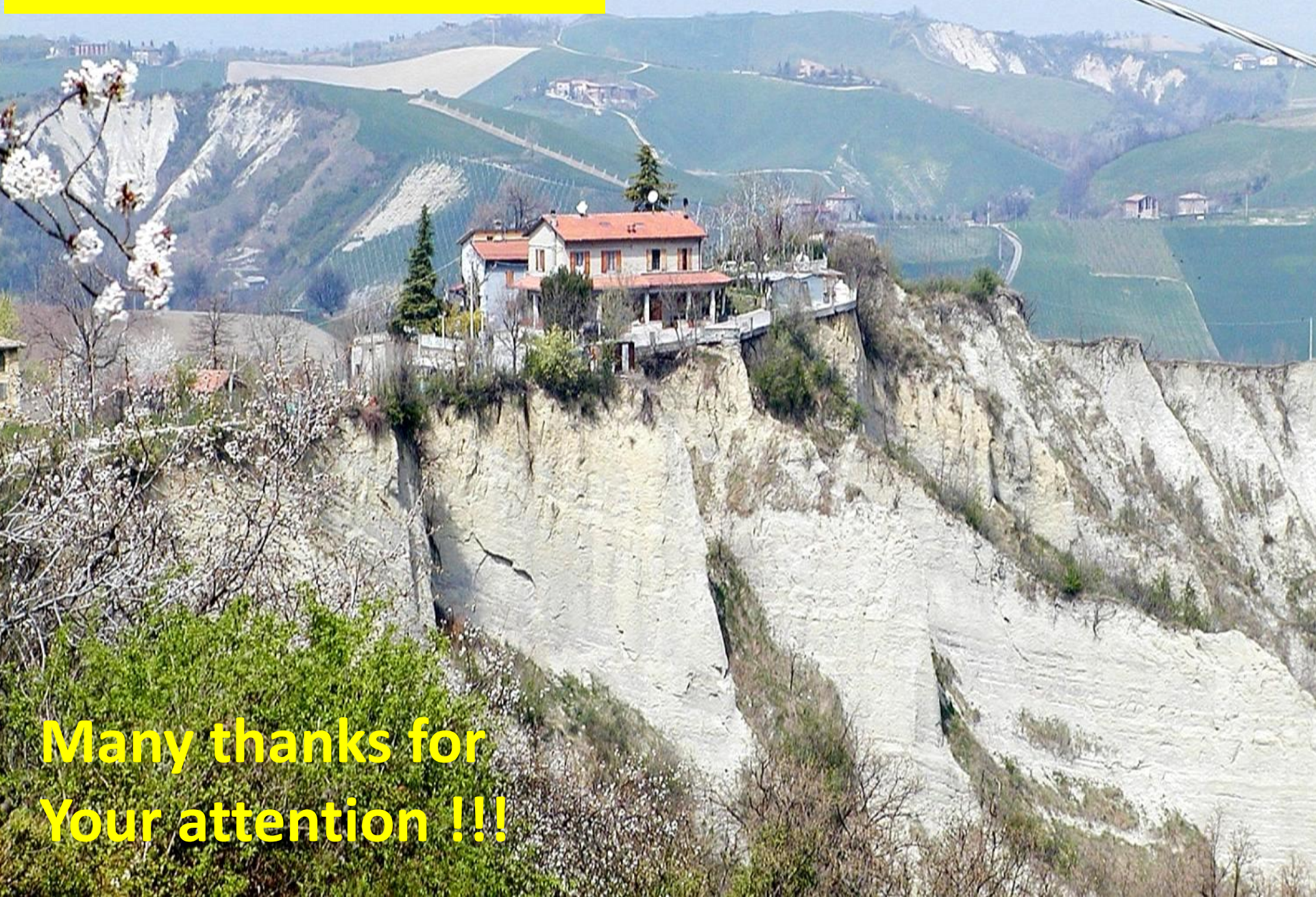
The IC model have a large set of potential applications such as **hot spot identification of primary sediment sources** to permanent drainage lines and **verification of impacts of eco-compatible mitigation measures to reduce or increase connectivity**. (without more complex Soil erosion models) (e.g. indications of Boardman, 2006)

IC can be easily transformed in a more physical based index (e.g. variants by Cavalli 2013 and Cassi 2010)

Potential application of IC to define **SDRL** can help to assessment of Sediment yield contribution due to Landslides and debris flow (PESERA-L model).



This presentation can be downloaded at  
**[www.lorenzo-borselli.eu](http://www.lorenzo-borselli.eu)**



**Many thanks for  
Your attention !!!**